

Original Research Article

Sedimentary structures and lithofacies found in a channel bar of Brahmaputra River in Panikhaiti, Kamrup District, Assam

ABSTRACT

The study deals with the recognition and interpretation of various bedform structures and lithofacies that had developed in a channel bar of the Brahmaputra River, in Panikhaiti near Guwahati, Kamrup district, Assam. The different bedform structures that are identified are small and mega ripples, water level cut marks, mud cracks, worm track and trails and rain drop imprint. Four vertical lithofacies sequences were also studied, which indicates the presence of eight lithofacies types (Miall, 1978). These are Trough cross bedded sand (St), Planar cross bedded sand (Sp), Horizontal laminated sand (Sh), Climbing ripple lamination (Sr), Convolute bedding (Fc), Flaser bedding (Sf), Massive sand (Sm) and Massive Mud (Fm). Based on the lithofacies associations, multiple floods are deciphered and their environmental significance has been interpreted.

Keywords: *Bedform structures, lithofacies, channel bar, Brahmaputra River, Panikhaiti, Kamrup District, Assam.*

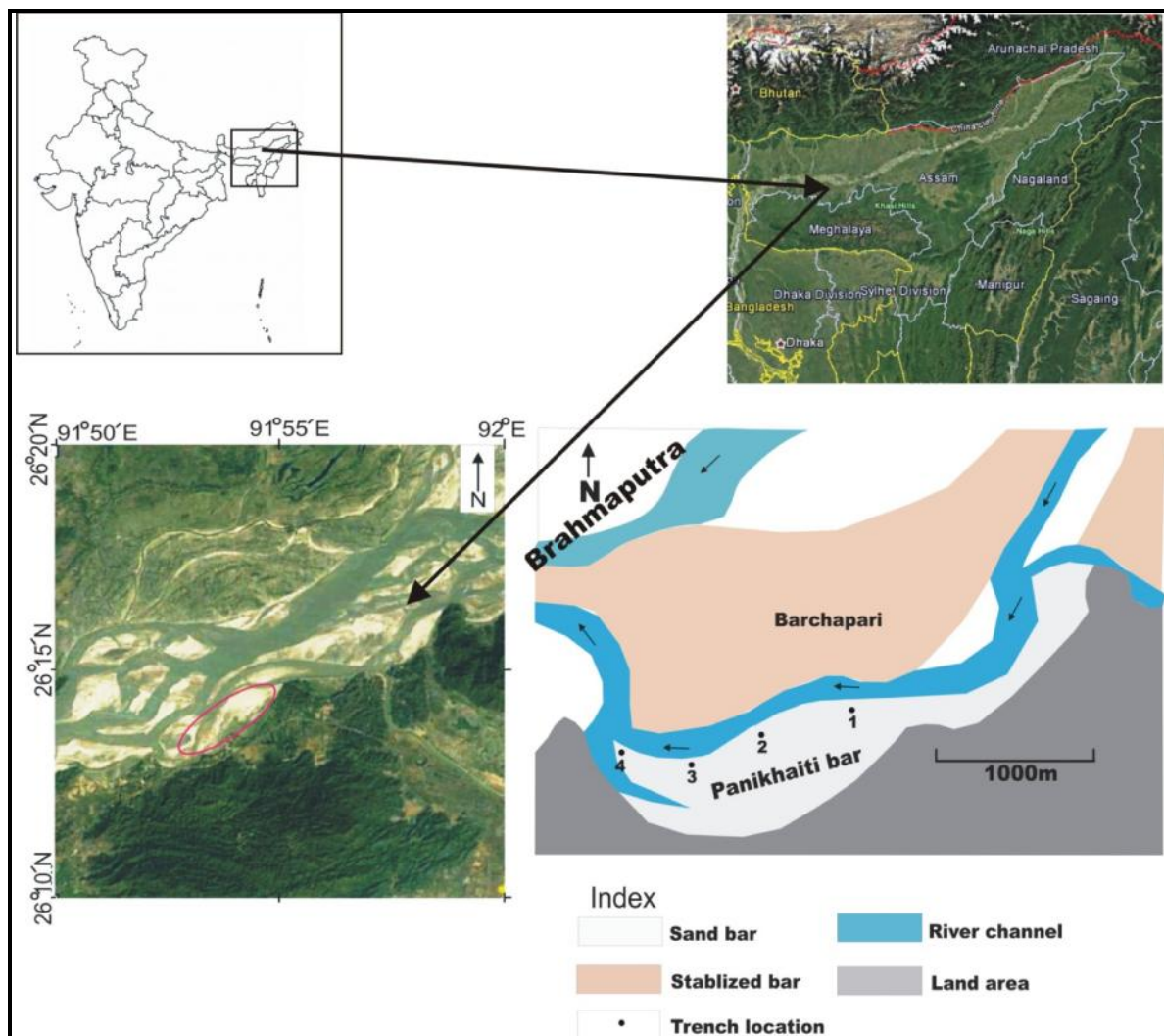
1. INTRODUCTION

The Brahmaputra River is a classic example of a braided river and is highly susceptible to channel migration and avulsion. It is one of the largest rivers in the world and is characterized by fine grained materials and channel bars with large seasonal discharge and sediment load in its lower reach (Reineck and Singh, 1980). Coleman (1969) described in detail, processes and various features of the Brahmaputra River in its lower reach in the deltaic plain of Bangladesh. Das and Borthakur (1995) presented a fluvial facies model of the Brahmaputra River near Nimatighat. Patgiri and Laskar (1996) worked on a channel bar near Guwahati and described different lithofacies types and sequences. Borkotoky (2015) published a paper on bedforms and lithofacies and interpretation of depositional environment of Brahmaputra River near Nemati, Jorhat, Assam.

The study area of the present bar is confined to parts of a bank-attached bar of the Brahmaputra River near Panikhaiti, Guwahati, Assam. This bank-attached bar is bounded within coordinates

32 $91^{\circ}52'20.97''$ and $91^{\circ}53'44.55''$ E and $26^{\circ}13'5.20''$ and $26^{\circ}13'39.14''$ N. The location of the bar is
 33 shown in Figure 1. The maximum length of the bar is about 2 km and average width is about 0.5 km.
 34 The bar is elongated in shape and acquires a pointed shape towards its downstream end.

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Fig. 1. Location map of the study area (Google Earth Image).

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2. METHODOLOGY

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The present study is concerned primarily with identification and recognition of various bedform
 40 structures and lithofacies developed on the channel bar, together with their environmental
 41 significance. Four vertical profile sections varying in depth from 1.7m to 2.25m are excavated. In order
 42 to study the three dimensional geometry of the lithofacies units, sections 1, 2, 3 and 4 are excavated
 43 both along and across the direction of the flow.

44

3. OBSERVATION AND DISCUSSION

45 **BEDFORM STRUCTURES**

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47 Various sedimentary structures have been observed on the bar surface. The structures are
48 mainly ripple marks, (small and mega ripple), water level marks, mud cracks, rain drop imprint, and
49 worm track and trails.

50 **Small ripples**

51
52 Most of the ripples are of lingoid type (Fig. 2A). They are discontinuous, isolated and are broken with
53 forward closures. Ripples range in length between 3 cm to 7 cm where as the height varies between 5
54 cm to 15 cm. They are produced as a result of the interaction of waves or currents on the sediment
55 surface (Reineck and Singh, 1980).

56 57 **Mega ripples**

58
59 Mega ripples are of larger height, greater length and asymmetrical crests and are undulatory in
60 nature. The wavelength of the mega ripples are greater than 4.5m, ripple index is found to be above
61 15 and the height varies from 35cm to 58cm and their crests are asymmetrical (Fig. 2B).

62 63 **Water level cut marks**

64
65 Water level cut marks are also called micro-terrace (Picard and High, 1973). They are usually
66 developed at the contact between water level and a sloping sediment surface and look like small
67 steps (Fig. 2C).

68 69 **Mud cracks**

70
71 Mud cracks form when mud is dewatered, shrinks and leaves a crack. Mud cracks of varied
72 dimensions are recorded in the area. They are polygonal in shape. Average thickness of individual
73 cracks range from 2.5 cm. to 30cm. while depth of the cracks varies from 0.5 cm to 65cm (Fig. 2D).

74 **Rain drops imprints**

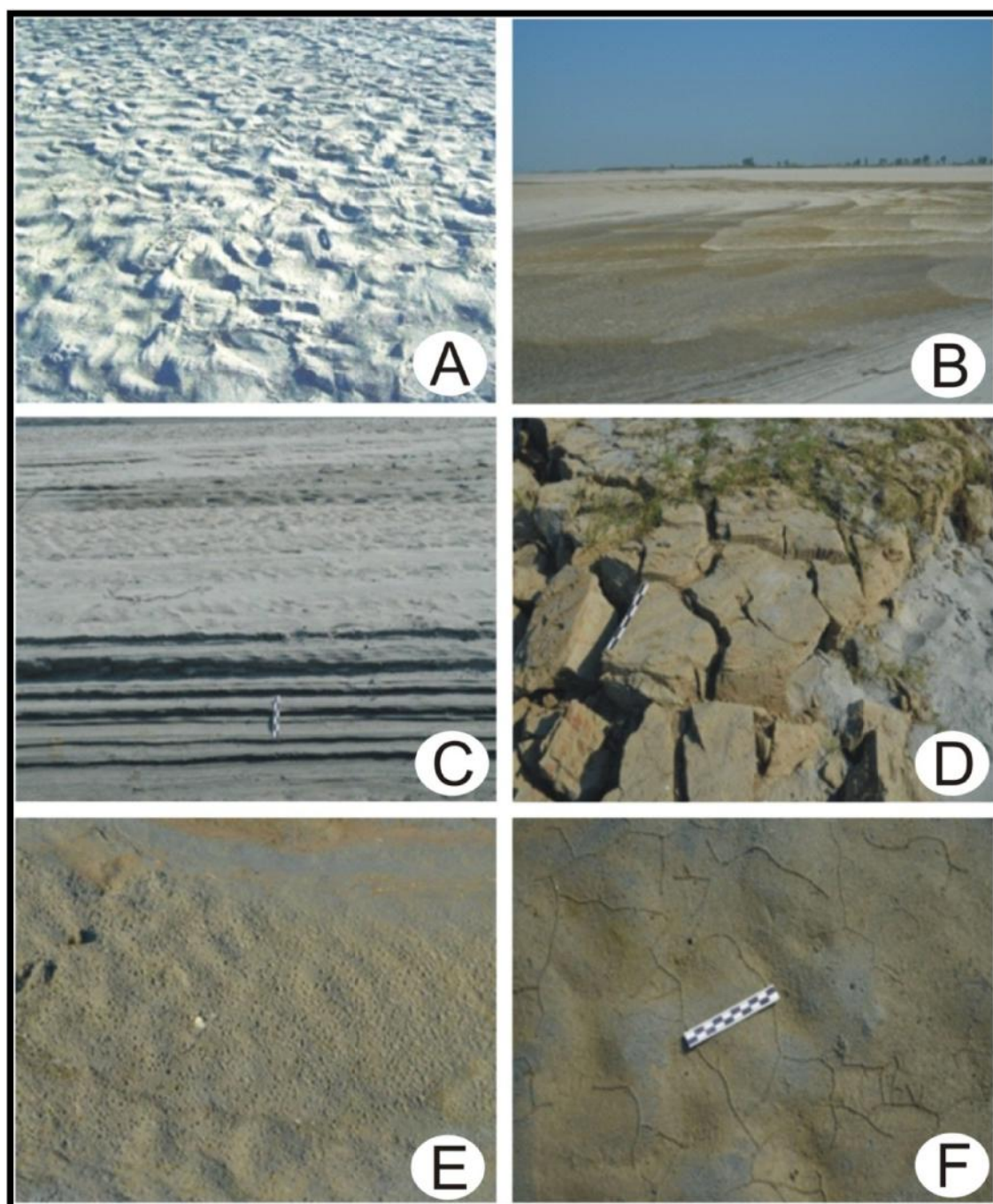
75
76 They are found on the top surface of loose sediment in the study area and are formed in wet mud or
77 loose fine sand and are circular (Fig. 2E).

78 79 **Worm tracks and trails**

80

81 They are produced by various organisms on the bar surface. Movement of organisms on the surface
82 of loose and soft sediment may cause the development of worm tracks and trails. (Fig. 2F).

83



84 Fig. 2. (A) Lingoid ripple. Current direction is from top right corner to down left corner, (B)
85 Mega ripples with undulatory crest line. Current direction from right to left, (C) Water level cut
86 marks. Scale is 30 cm, (D) Mud cracks, scale is 30cm, (E) Rain drop Imprints, (F) Worm tracks
87 and trails, scale is 10 cm.
88

89 LITHOFACIES ANALYSIS

90
91 The term "Lithofacies" is used to describe various depositional units with from the point of view of their
92 characteristic internal structures, grain size and other physical properties. Different structures in a
93 sedimentation unit may develop either due to variations in the energy condition of the depositional
94 environment or due to variation of grain size of the available material for sedimentation in a constant
95 energy condition (Fritz and Moore, 1988).

96
97 Based on sedimentary structures 8 (eight) lithofacies varieties are recognized and coded following
98 Miall (1978, 1996), and are described below.

99

- 100 1. Trough cross bedded sand (St)
- 101 2. Horizontal laminated sand (Sh)
- 102 3. Planar cross bedded sand (Sp)
- 103 4. Climbing ripple laminated sand (Sr)
- 104 5. Convolute bedding (Fc)
- 105 6. Flaser bedding (Sf)
- 106 7. Massive sand (Sm)
- 107 8. Massive mud (Fm)

108

109 1. Trough cross bedded sand (St)

110

111 A cross-bedded unit where bounding surface is curved and the units are trough shaped is called
112 trough cross-bedded (Reineck and Singh, 1980). This facies consist of fine to very fine sand. It is
113 further sub-divided into two types St₁ and St₂ on the basis of their dimension. Facies St₁ represent
114 trough cross bedding with set thickness greater than 4 cm, while St₂ represents trough cross bedding
115 with set thickness less than 4 cm. The thickness of the facies units vary between 15 cm and 80 cm
116 (Fig. 3A). The trough cross bedded sand indicates their deposition under increasing flow strength in
117 lower flow regime conditions.

118 2. Horizontal Laminated Sand (Sh)

119

120 This lithofacies is composed of very fine sand (Fig. 3B) and range in thickness between 8 cm and 70
121 cm. This facies represents deposition under high velocity condition in the plane bed phase of upper
122 flow regime (Parkash et. al. 1983).

3. Planar Cross bedded sand (S_p):

This facies is composed of medium sand and the beds are tabular to wedge-shaped. The foreset angle varies from 25° to 45° . The thickness of the beds varies from 7 cm to 30 cm (Fig. 3C). Planar cross-bedded sand forms due to the migration of straight crested large or small ripples on the surface of the bar supplying sand, which has slipped down and was deposited on the avalanche face of the bar (Singh and Bhardwaj, 1991).

4. Climbing –Ripple Laminated Sand (S_r)

Climbing ripple lamination are the internal structure formed in non cohesive materials from migration and simultaneous upward growth of ripple produced either by current or waves. Climbing ripple laminations, particularly Type-II (drift type) (Reineck and Singh, 1980) are found in abundance in the study area. They indicate a relatively high rate of sediment deposition in comparison to transportation (Mckee, 1966). They range in thickness from 20cm to 30 cm (Fig. 3D).

5. Flaser bedding (S_f)

Ripple bedding in which mud streaks are preserved completely in troughs and partly in crests is known as Flaser bedding (Reineck and Singh, 1980) (Fig. 3E). The thickness of the bed varies from 5 cm to 25 cm.

6. Convolute bedding (F_c)

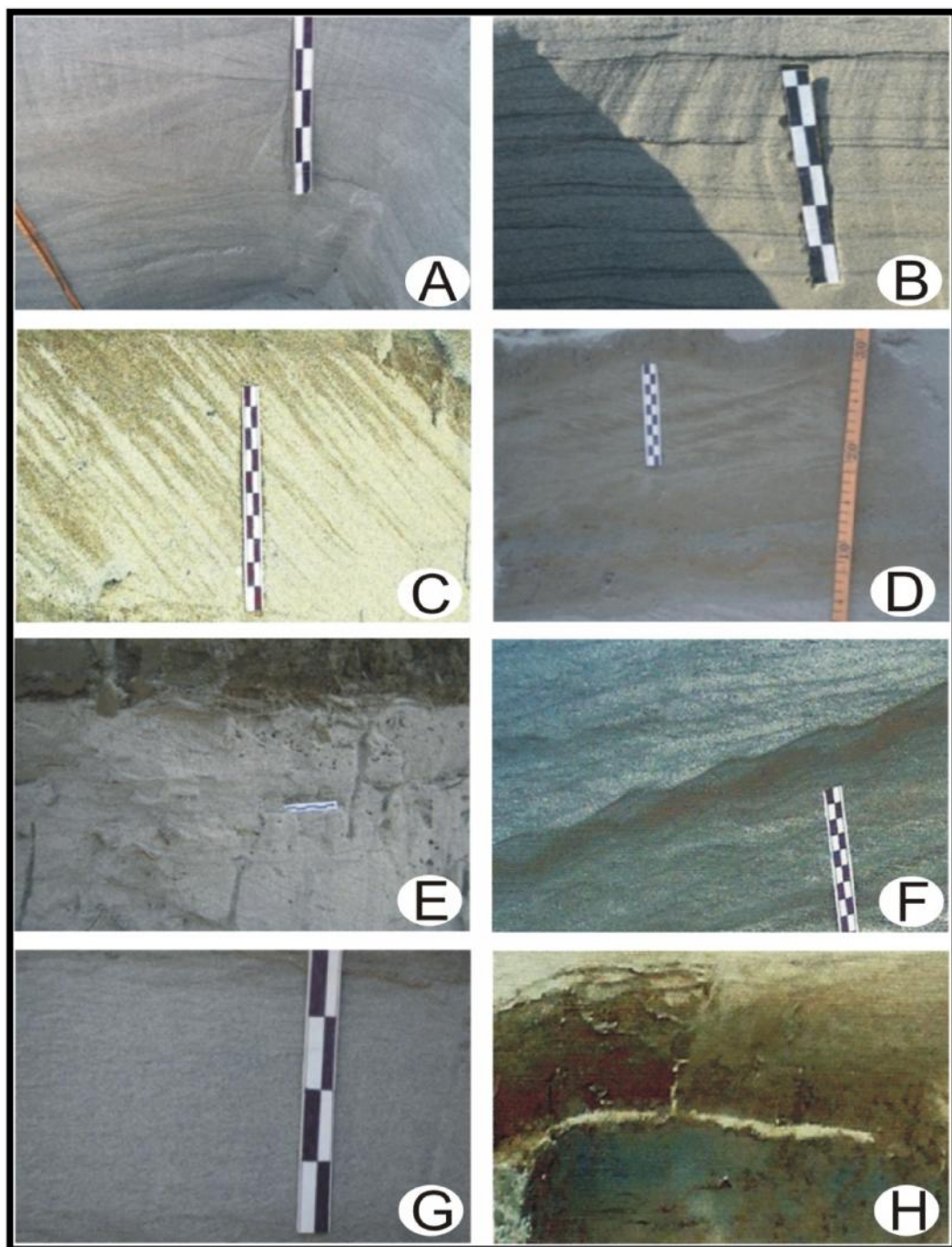
Convolute bedding is a structure showing marked crumpling or complicated folding of laminae of a well defined sedimentation unit (Kuenen, 1953; Potter and Pettijohn, 1963). This facies is well developed in very fine sand. The convolute are asymmetry and the axial plane inclined to the down current direction. The thickness of the bed is 18 cm (Fig. 3F).

7. Massive sand (S_m)

This lithofacies is characterized by massive nature and does not show any internal stratification and average thickness of strata is about 15 cm (Fig. 3G). The massive units are possibly post-depositional products of local nature formed by underwater slumping in still-soft sediment (Ray, 1976).

8. Massive Mud (F_m)

159 Facies Fm is composed of homogeneous massive mud layers composed of variable amounts of silt
 160 and clay. The thickness of the beds varies from 5 cm to 30 cm. They are formed during the waning
 161 stages of flood by vertical accretion of fine suspended sediments under quiet environment (Fig. 3H).



162 Fig: 3 (A) Trough cross bedded sand, (B) Horizontal laminated sand, (C) Planar cross
 163 bedded sand, (D) Climbing ripple laminated sand, (E) Flaser bedding, (F) Convolute
 164 bedding, (G) Massive sand and (H) Massive mud.
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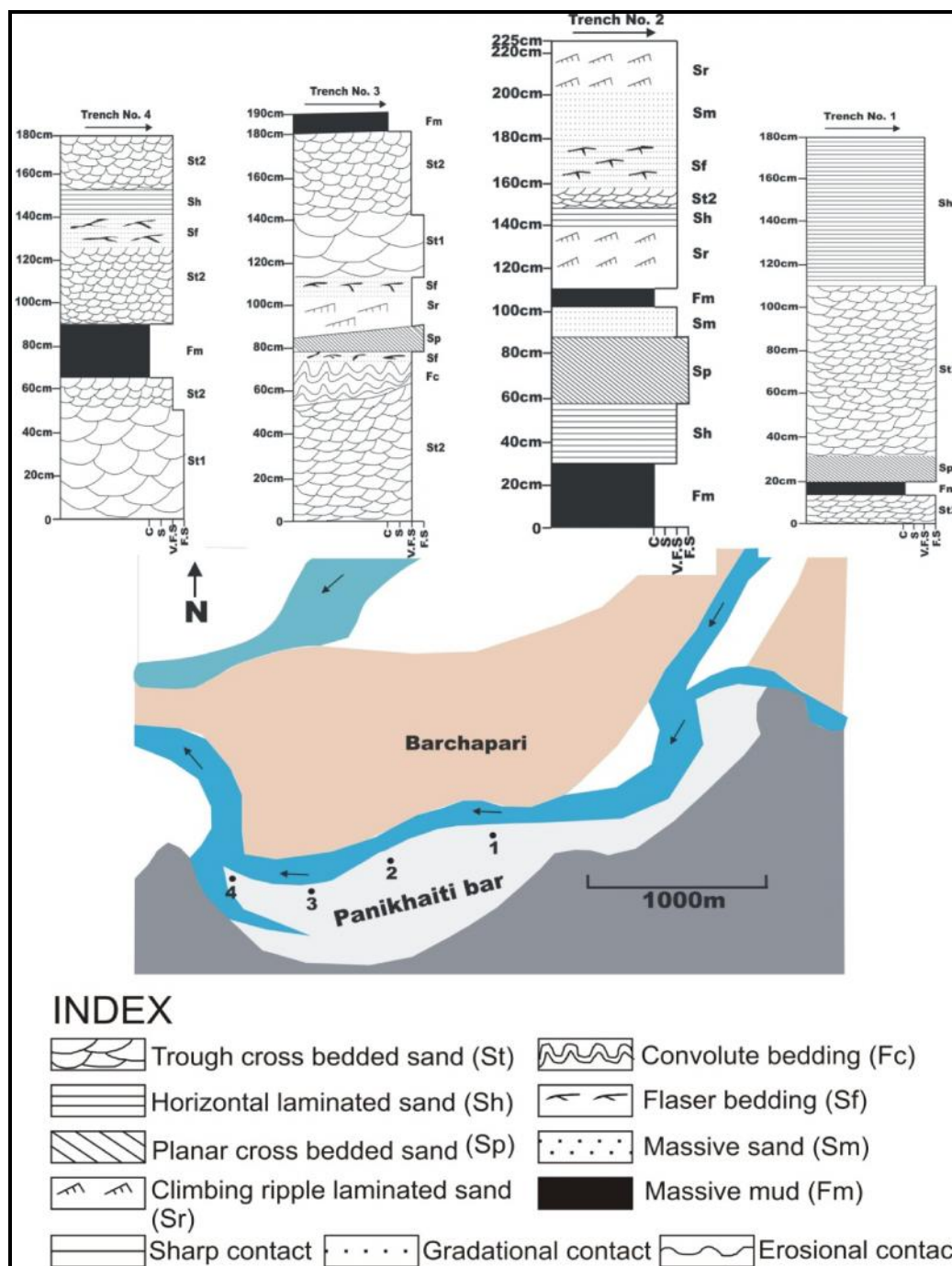


Fig. 4. Vertical profile section showing distribution and association of various lithofacies types.

LITHOFACIES ASSOCIATION

Fig. 4 shows the vertical and lateral distribution of the lithofacies of the bar. **Trench 1** represents a 180 cm thick sequence from the upstream portion of the bar. The lower portion of the sequence is composed of small scale trough cross bedded sand (St_2). This facies is overlain by massive mud (Fm) which marks the end of a flood cycle. Facies Fm is overlain by planar cross bedded sand (Sp). Again Sp is overlain by small scale trough cross bedded sand (St_2) facies. The foreset laminae angle of facies Sp varies from 30° to 35°. They are deposited by migration of large two dimension ripples in lower flow regime condition. Facies St_2 is overlain by horizontal laminated sand (Sh) which develops under upper flow regime conditions.

The facies association reflects a gradual transition from lower flow regime conditions to an upper flow regime condition in the upstream portion of the bar.

Trench 2 represents the middle portion of the bar where the sedimentary sequence becomes 225 cm thick. The lower portion of the sequence contains massive mud (Fm). This facies is followed by horizontal laminated sand (Sh) which represents deposition under upper flow regime conditions. Facies Sh is overlain by planar cross bedded sand (Sp). Sp is in turn overlain by massive sand (Sm) followed by Fm. These facies associations represent a complete flood cycle. Climbing ripple laminated sands (Sr) occur above Fm. Facies Sr is subsequently overlain by Sh which indicates a change from lower flow regime to upper flow regime condition. A drop in flow velocity is indicated by the occurrence of Facies St_2 above Facies Sh. Flaser bedding (Sf) is found above Facies St_2 . Facies Sf is subsequently overlain by massive sand (Sm) and Facies Sr, which indicates prevalence of lower flow regime condition during the deposition of the most recent sediments.

Trench 3 represents the middle portion of the bar where the sequence becomes 190 cm thick. The lower portion of the sequence is composed of small scale trough cross bedded sand (St) which is followed upwards by convolute bedding (Fc) and Flaser bedding (Sf). Facies Sf is overlain by planar cross bedded sand (Sp). They indicate a lower floor regime condition. Facies Sp is further overlain by climbing ripple laminated sand (Sr) and flaser bedding (Sf). This facies is followed by large scale trough cross bedded sand (St_1). St_1 is overlain by small scale trough cross bedded sand (St) and finally by massive mud (Fm) unit. A complete flood cycle is also found in this sequence.

Trench 4 is located in the downstream portion of the bar and is 180 cm in thickness. The lower portion of the sequence is composed of large scale trough cross bedded sand (St_1) and is followed upwards by small scale trough cross bedded sand (St_2) and massive mud (Fm). Facies Fm is overlain by facies St_2 followed by flaser bedding (Sf). Sf is overlain by horizontal laminated sand (Sh) which indicates an upper flow regime condition. This facies is overlain by facies St_2 and indicates a drop in current velocity during their deposition.

4. CONCLUSIONS

- i. Various bedform structures are found on a bar. In case of the present bar, they are represented by small and mega ripples, water level cut marks, mud cracks, rain drop imprints and worm tracks and tails.
- ii. Eight different types of lithofacies could also be identified in the bar. They are Trough cross bedded sand (St), Horizontal laminated sand (Sh), Planar cross-bedded sand (Sp), Climbing ripple laminated sand (Sr), Convolute bedding (Fc), Flaser bedding (Sf), Massive sand (Sm) and Massive mud (Fm). They represent deposition under both lower and upper flow regime conditions.
- iii. The lithofacies associations show development of both complete as well as incomplete fining-upward flood sequences. Incomplete flood sequence show association of fewer lithofacies types and indicate the simultaneous erosion and deposition by floods.

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