

HYDRO-PHYSICAL PROPERTIES AND MICRONUTRIENT STATUS OF MAYO-GWOI FLOODPLAIN SOILS, IN TARABA STATE, NIGERIA

Abstract

The aim of this research was to study the hydro-physical characteristics and the status of micronutrients in floodplain soils of Mayo-Gwoi in Taraba state, Nigeria. A detailed soil survey was conducted at the Mayo-Gwoi floodplain using rigid grid approach. Observations were made at 100m regular intervals. Two profile pits were dug and sampled from each of the three soil mapping units identified. The samples were analyzed and characterized as follows; texture vary from loamy to sandy clay loam, bulk density (1.30 g/cm^3), particle density (2.63 g/cm^3), water holding capacity (37.5%), water content at field capacity (36 %) and wilting point (24 %), zinc (1.0 Mg/Kg), iron (6.3 Mg/Kg), copper (0.5 Mg/Kg) and manganese (6.3 Mg/Kg). These soils showed some evidence of degradation and could be productive if subjected to appropriate management and maintenance.

Key words: Micronutrients, physical characteristics, Mayo-Gwoi floodplain soils, rigid grid

Introduction

Knowledge of soil physical and chemical properties are key to making agronomic and environmental decisions. However, the heterogeneous nature of these soil properties mainly due to changes in structural and chemical composition of soil minerals from microscale to global scale (Brady, 2006), makes this decision difficult. The dynamics of soil moisture including the water retention capacity and soil nutrient status determine to a large extent the soil productivity. According to Havlin (2003), micronutrients are as important in plant nutrition as macronutrients; though they simply occur in plants and soils in much lower concentrations. It has been observed that plants grown in micronutrient-deficient soils exhibit similar reductions in productivity as those grown in macronutrient-deficient soils. Micronutrients in soils exist in the form of elements in primary and secondary minerals; adsorbed to mineral and organic matter surfaces; incorporated in organic matter and microorganisms; incorporated into solution, depending on the source of the micronutrients. Understanding the relationships and dynamics among these forms is essential for optimizing plant productivity in micronutrients deficient soils. The availability of micronutrients to plants is determined by both the total amount of the nutrient in the soil and the soil's

31 properties. Other factors such as crop species and variety can also influence the degree to
32 which micronutrient levels affect crop production.

33 The hydro-physical properties of soil including the soil moisture dynamics and water
34 retention capacity are influenced by land use. Land use is defined by Vink (1975) as any kind
35 of permanent or cyclic human intervention to satisfy human needs from a complex of natural
36 and artificial resources which together constitute land. Therefore, at any given time, land use
37 is resultant of interplay of available land resources with cultural, social and economic
38 conditions of the past and present developments. According to Odingo (1991), land use in
39 Africa is excessively influenced by the patterns of use which have been set by previous
40 generations and administrations depending on the kind and level of technology available to
41 them at the time. Ojanuga (1991) writes that land use pattern is an expression of the
42 agricultural system, representing the interaction between the ecological and economic factors
43 in which the vital link is the decision maker, namely, the farmer.

44 It is of concern that in spite of the increasing interest in fadama farming in Taraba State, there
45 is no available literature on the hydro-physical properties and micronutrient levels of the
46 floodplain soils. Thus, based on the realization that such information forms the background to
47 an efficient and judicious use of the soil resources, the objective of this study was to study the
48 hydro-physical properties and the status of micronutrients in the soils of Mayo-gwoi
49 floodplain.

50

51 **Materials and Methods**

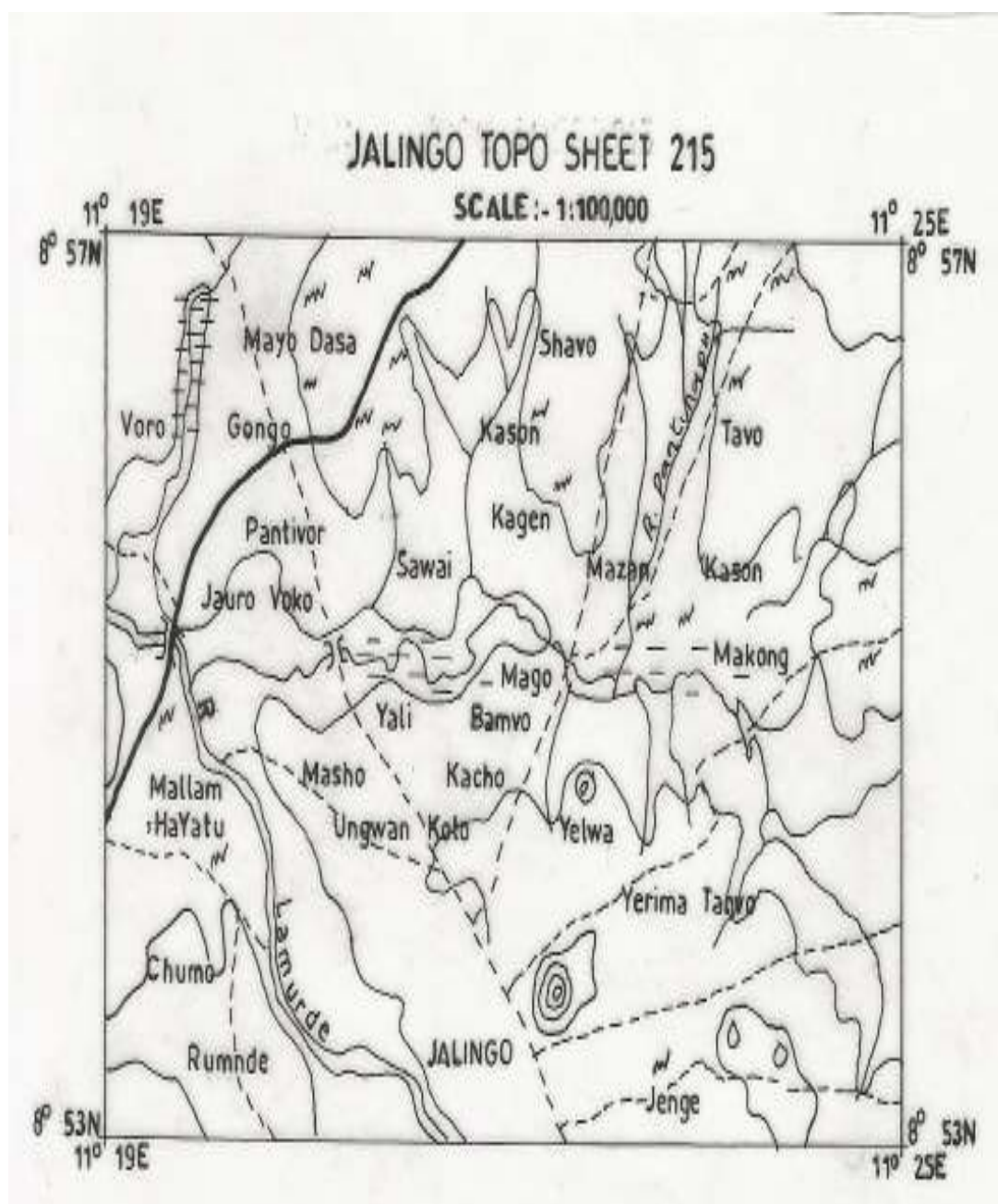
52 **Description of the study area**

53 The Mayo-Gwoi floodplain is located between latitudes $8^{\circ} 53'$ and $8^{\circ} 85'$ North and
54 longitudes $11^{\circ} 23'$ and $11^{\circ} 75'$ East. It covers an area of about 120ha. It's located in Jalingo
55 city, the capital of Taraba State. The geology of the area has been described by Bawden
56 (1972) as a basement complex and the rocks are mainly pre-cambrian granitic and migmatite
57 gneisses. Jalingo lies within the tropical hinterland climate region. The region is characterized
58 by double maxima rainfall pattern which has about four months of dry season with relative
59 humidity having generally over 80% in the morning and falls to between 50 and 79% in the
60 afternoon. The dry and wet seasons are controlled by the annual migration of the intertropical
61 zone of convergence (ITZC). The dry season is characterized by the dry dust, laden with
62 harmattan winds coming across the Sahara desert and occurring between November and
63 February of every year. The wet season sets in by April and lasts until October (Iloeje, 1981).

64

Fig.

1.



Sampling and Laboratory Analysis

Detailed soil survey was conducted at the Mayo-Gwoi floodplain soils which is the study site. A Rigid grid method was employed for surveying the area, baseline was drawn and traverses perpendicular to the baseline were cut and observations were made at 100m regular intervals. Based on the information obtained in the field, three (3) mapping units were identified viz mapping units 1, 2 and 3 respectively. In mapping units 1 and 3, two profile pits each were sunk while mapping unit two had three profile pits making seven profile pits. Profile pits

74 were sampled according to the pedogenetic horizons. Soil colour was determined using
75 Munsell soil colour chart. Particle size distribution was determined by Bouycous hydrometer
76 method using sodium hexametaphosphate (Calgon) as the dispersant and the textural class
77 determination adopted was the USDA textural triangle method Jaiswal (2003). The bulk
78 density was by core method (Blake and Hartge, 1986). The particle density was determined
79 as described by Jaiswal, (2003). Soil moisture content was determined using the gravimetric
80 method as described by Jaiswal (2003) while the water holding capacity of the soil was
81 determined by the method of Klute (1986).The porosity of the soil was determined according
82 to the method described by (Blake and Hartge, 1986) in Jaiswal (2003).

83

84 Results and Discussions

85 **Table 1: Physical properties of Mayo-Gwoi flood plain soils**

Mapping Unit	Profile Depth	B.D g/cm ³	PD g/cm ³	Porosity (%)	F.C (%)	WP (%)	AV H ₂ O (mm)	WHC (%)	Sand (%)	Silt (%)	Clay (%)	Textural class
MU1	0-18	1.42	2.65	46	25	14	11	37.1	56	21.2	22.8	Scl
	18-61	1.44	2.64	46	23	14	9	37.2	66	11.2	22.8	Scl
	61-82	1.41	2.64	47	25	16	9	37.2	62	13.2	26.8	Scl
	82-99	1.26	2.63	52	40	29	11	38.1	52.8	38	22.8	S
	99-144	1.35	2.62	45	22	14	8	32.2	54	5.2	40.8	Scl
	144-158	1.33	2.64	50	31	16	15	38.2	28	43.2	28.8	Cl
	158-178	1.24	2.64	53	37	28	9	32.2	47.2	40	52.8	Sl
	178-200	1.22	2.65	54	43	27	16	33.3	52.0	43.2	46.8	Sl
	0-14	1.32	2.63	50	32	23	9	40.2	10	61.2	28.8	Scl
	14-41	1.34	2.64	49	30	21	9	40.3	58	3.2	38.8	Sc
	41-78	1.49	2.64	44	21	13	8	40.2	58	19.2	22.8	Scl
	78-122	1.33	2.65	50	31	19	12	40.1	78	3.2	18.8	Sl
	122-141	1.4	2.63	47	26	15	11	39.2	74	3.2	22.8	Scl
	141-180	1.34	2.59	49	30	21	9	34.6	58	3.2	38.8	SC
	180-200	1.42	2.62	46	24	14	10	35.7	52	5.2	42.8	Scl
MU2	0-30	1.23	2.64	54	42	26	16	38.2	42	31.2	26.8	L
	30-50	1.29	2.65	51	34	15	18	38.4	8.0	29.2	62.8	C
	50-85	1.35	2.65	49	29	15	14	38.5	58	19.2	22.8	Scl
	85-140	1.32	2.59	50	32	18	14	33.3	35.2	58.0	26.8	L
	140-175	1.18	2.64	56	51	38	14	34.2	10	45.2	44.8	SL
	0-38	1.24	2.64	53	41	24	16	39.6	32	35.2	32.8	Cl
	38-62	1.33	2.64	50	31	19	12	33.4	10	47.2	42.8	Sc
	62-120	1.24	2.65	53	40	23	17	40.6	42	23.2	34.8	Cl
MU3	0-18	1.24	2.64	53	41	24	16	39.6	32	35.2	32.8	Cl
	18-54	1.22	2.64	54	42	26	17	39.4	58	21.2	20.8	Scl
	54-105	1.29	2.64	51	36	24	12	39.7	2.0	53.2	44.8	Sl
	105-134	1.22	2.63	54	42	26	17	35.6	8.0	51.2	40.8	Sc
	134-184	1.44	2.64	46	24	13	10	34.7	4.0	19.2	22.8	Sc
	0-34	1.18	2.57	56	54	42	12	40.3	36	21.2	42.8	C
	34-52	1.22	2.61	54	46	34	12	40.5	22	5.2	72.8	C
	52-93	1.18	2.58	55	53	42	11	33.4	18	9.2	72.8	C
	93-130	1.25	2.63	53	41	31	10	35.4	40	3.2	56.8	C

	0-38	1.25	2.64	53	40	28	12	41.2	32	17.2	50.8	C
	38-60	1.24	2.61	53	43	32	11	40.6	34	9.2	56.8	C
	60-90	1.37	2.57	48	43	32	15	40.2	32	45.2	22.8	L
	90-120	1.25	2.64	53	41	29	12	40.4	32	16.8	51.2	C
	120-150	1.37	2.63	48	43	32	15	35.3	32	45.2	22.8	L
		1.18-	2.57-					32.2-		3.2-	18.8-	
Range	-	1.49	2.65	44-56	21-54	13-42	8-18	41.2	2-78	61.2	72.8	
Mean	-	1.30	2.63	51	36	24	12	37.5	39	25.9	37.8	

86 BD= bulk density; PD =particle density; FC =field capacity; WP =wilting point; AV.H₂O =available
87 water; WHC= water holding capacity, C=Clay, SC=Sandy clay, Cl=Clay loam, Scl=Sandy clay loam,
88 L=Loam, S=Sand

89

90 **Table 2.** Micronutrients in Soils of Mayo-Gwoi Floodplain

Pedon Number	Pedons Depth (cm)	Zn (Mg/kg)	Fe (Mg/kg)	Cu (Mg/kg)	Mn (Mg/kg)
P1	0-18	1.1	9.7	ND	6.4
	18-61	2.6	8.1	0.5	4.4
	61-82	2.7	7.4	0.6	20.8
	82-99	2.1	5.6	ND	0.8
	99-144	2.6	3.1	0.5	4.6
	144-158	0.4	2.7	0.5	4.9
	158-178	0.5	2.8	0.5	10.2
	178-200	0.7	2.2	0.4	7.8
P2	0-14	1.5	8.1	0.5	5.9
	14-41	1.6	8.8	0.5	5.6
	41-78	1.6	5.4	0.5	9.2
	78-127	2.6	8.9	0.3	11.7
	127-141	2.4	5.6	0.4	3.1
	141-180	0.5	3.9	0.3	6.9
	180-200	0.3	7.7	0.7	4.0
P3	0-30	1.1	8.7	0.4	15.1
	30-50	0.5	6.9	0.4	6.3
	50-85	0.5	5.1	0.7	3.9
	85-140	0.8	7.8	0.5	7.1
	140-175	0.6	5.0	0.4	8.7
P3	0-32	0.5	6.2	0.6	2.4
	32-62	0.8	5.7	ND	6.1
	62-120	0.8	5.7	0.5	6.1
P4	0-18	0.3	5.6	0.3	2.8
	18-54	0.3	7.7	0.4	3.6
	54-105	0.5	6.0	0.4	6.9
	105-134	0.2	5.1	0.5	1.1
	134-184	0.6	5.0	0.6	7.2
P5	0-38	0.8	9.7	0.5	8.0
	38-52	0.7	5.9	0.5	6.3

	52-93	0.8	4.0	0.7	8.9
	93-130	0.6	7.1	0.4	9.2
P6	0-38	0.5	9.0	0.5	5.3
	38-60	0.6	5.7	0.6	5.4
	60-90	0.7	8.0	0.5	7.0
	90-120	0.7	8.7	ND	5.9
	120-150	0.1	3.4	0.3	6.4
Range	-	0.1-2.7	2.2-9.7	0.3-0.7	0.8-20.8
Mean	-	1.0	6.3	0.5	6.7

91 Note: ND = No data

92 Soil textures were variable within the mapping units, surface texture ranged from loamy sand
93 through sandy clay loam to clay. The sand contents of the profiles changed with depth, it has
94 clay texture in both surface and subsurface horizons, these variations may be due to
95 differences in parent materials and topography (Brady and Weil, 2002). The clay content of
96 the soils was low while the sand content was high. Soil structures were also variable: being
97 weakly developed in pedons that are moderately coarse-textured (pedons 1-5) and well
98 developed in a pedon with high clay content. This lack of structural development in the
99 horizons could be attributed to the effect of low water table (Udo, 2001). The bulk density of
100 the soils ranged from 1.22 g/cm³ to 1.44g/cm³ (Table 1). These values are considered safe for
101 root penetration since root penetration may be hindered in soils having bulk density above
102 1.75 g/cm³ (Esu, 2005).

103 The water holding capacity in both the surface and subsurface horizons of the soils
104 were low (Table 2) and this could be attributed to the low organic matter content. Organic
105 matter has been known for its importance at improving water retention capacity of most
106 surface soils (Brady and Weil, 2002). Consequently, water at field capacity, water at wilting
107 point and available water in the soils are low but fall within the range that cannot cause any
108 negative effect to most arable crops (Brady, 2002). The concentration or content of zinc in
109 these soils were low (Table 2) and could be due to low organic matter content of the soils
110 (Brady and Weil, 2002).The iron content of the soils were generally low to moderate (Table
111 2). The low iron content could be due to transformation and redox reactions (Brady, 2006).
112 Similar result was reported by Mustapha *et al*, (2003) on the study of profile distribution of
113 some Hydromorphic soils of Dass, Bauchi State. However, the level of available copper in
114 the soils was deficient in all the pedons (Table 2) at both the surface and subsurface horizons
115 which could be attributed to the low crystal concentration of copper in the soils dynamics
116 (Havlin *et al*, 2003).The level of manganese was moderate at the surface and subsurface
117 horizons in all the pedons (Table 2). This could also be attributed to the low organic matter

content and the acidic nature of the soils shown by their pH values. Brady and Weil (2002) and Havlin *et al*, (2003) opined that the availability of most of the micronutrients in soils depend on the soil pH and organic carbon contents. Consequently, improving the organic carbon contents of the soil which serves as the mainstay of most extractable soil micronutrients (Brady and Weil, 2006), could help to improve the productivity of Mayo-Gwoi flood plain soils

CONCLUSION

The study highlights the hydro-physical characteristics and the micronutrient status of soils from Mayo-Gwoi floodplain in Taraba state, Nigeria. The soils are characterized to have varying textures ranging from loam through loamy sand to sandy clay loam, low to moderate bulk density values as well as water holding capacity. The soils also showed low to moderate levels of the micronutrients (zinc, iron, copper and manganese). Good productivity of these soils would be ensured by adequate soil management strategies such as addition of organic matter in the form of poultry manure to improve the soil since organic matter is a major source of micronutrients in soils as well as improving soil physical health.

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