# **Original Research Article**

## Categorization of Morphometric Surface through Morphometric Diversity Analysis in Kushkarani River Basin of Eastern India

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### 6 ABSTRACT

7 The present work attempts to build up and categorize morphometric surface of distinct identities and 8 potentialities considering 16 morphometric parameters as spatial data layers in GIS environment. For 9 constructing the morphometric diversity model relative importance of the selected parameters is 10 assessed and weighted compositing is being made. Separately, relief and drainage diversity models 11 have been prepared and finally entire surface is categorized into some distinct morphometric sub 12 units having their own terrain characteristics and associated cultural features. Out of total river basin 13 only 11% area comes under highly diversified surface and maximum part of the area comes under 14 quite monotonous surface with least variation. Highly diversified surface is prone to rilling, gullying and 15 consequently soil erosion. It is also tried to identify whether relief and drainage model exert significant 16 impact on morphometric diversity model. It is identified that relief diversity model carries significant 17 impact on morphometric diversity model as indicated by high correlation coefficient (r=0.94).

Keywords: Mophometric surface, Weighting of parameters, Weighted linear combination (WLC),
 Drainage diversity model (DDM), Relief diversity model (RDM) and Morphometric diversity model
 (MDM).

#### 21 **1. INTRODUCTION**

Landscape is an essentially visual phenomenon or as a particular constitution of topography, land 22 use, vegetation cover and settlement pattern <sup>1,2</sup> and morphometric surface is one of the fundamental 23 terrain base on which cultural and ecological conditions can be draped. Morphometric analysis 24 25 includes a good number of quantitative indicators based on which terrain units can be recognized and 26 each such terrain unit provides a strong base for analyzing terrain processes and associated cultural fabrics <sup>3,4</sup>. Previous studies focused on morphometric characteristics of various drainage basins 27 using GIS and remote sensing technique for terrain analysis <sup>5-11</sup>. The Morphometric parameters 28 29 calculated in the study are flow directions, flow accumulation, upstream and downstream, stream link, stream network, stream order, and digital elevation models (DEMs) etc.<sup>12,13</sup>. Generally in this studies 30 31 emphasize segregated statistical analysis or mapping of the individual morphometric parameters. Among then some studies correlate morphometric parameters and potential ground water 32 investigation <sup>14-20</sup>, recharge estimation<sup>21-23</sup> and surface water availability<sup>24-26</sup> etc. Some studies laid 33 importance on sub basin priority for promising planning unit <sup>27-30</sup>. The present work pursues how to 34 35 indentify morphometric surface units using different morphometric parameters as spatial data layers. 36 Primarily, drainage and relief morphometric surfaces are recognized and finally morphometric surface is also being identified. Here attempt is also being made regarding prioritizing incorporated
 parameters considering their importance towards bringing morphometric diversity.

39 Morphometric diversity analysis actually helps to find out some nearly homogenous landscape units 40 which are uniform in morphometric characteristics. Each unit can be considered as a homogenous 41 hydro-geomorphic spatial unit with unique nature and potential viability of land use. Landscape unit 42 with highly diversified morphometric characteristics is healthy in geomorphic processes and forms and 43 active in terms of ongoing terrain processes. On the contrary, lowly diversified landscape unit is 44 characterized by relatively monotonous processes and forms. Each landscape unit is uniquely 45 potential for resources and economic activities. This fact becomes highly vivid when range of 46 morphometric differences is very high and size scale is very large. In case of small basin over 47 homogenous physiographic unit, it is not so clear and therefore, not so effective for dictating the LULC characteristics <sup>31, 32</sup>. But some indirect influences can be streamlined. For example, morphometric 48 49 diversity unit has its potentiality to characterize soil loss processes, runoff processes, recharge 50 activities etc. These are often responsible for controlling LULC and livelihood opportunities. 51 Hypsometric surface is well explained in relation to geology of the present study area using hypsometric integral (HI) of different sub basins by Khatun and Pal<sup>33</sup>. This surface envisages the 52 53 erosion processes, soil loss, sediment deposition pattern etc.

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#### 2. STUDY AREA

56 Kushkarani river (length: 35 km) basin, covering an 172 sq km area (see fig-1), is a sub basin of 57 Mayurakshi river system located mainly over the western part of Chottonagpur plateau fringe in 58 Birbhum district of West Bengal and Jamtara district of Jharkhand with 23°54' 36" N to 24'00' N 59 latitudes and 87°14'24" E to 87°30' E longitudes. The total basin area comes under rarh tract topography <sup>34</sup> with lateritic soil formation <sup>35</sup> which is mainly received by flowing rivers from 60 61 Chottonagpur plateau. This river pours into Tilpara barrage which is located over master stream Mayurakshi river <sup>36</sup>. The elevation of this catchment varies from 155 m. (at the source region) to 62 62 63 m. (at the confluence region). Maximum area of the basin <mark>shows</mark> rugged topography with an average 64 elevation of 108 m. Average slope of the basin is 1-4 degree whereas it is <1 degree in the confluence part of the basin measured as per Wentworth's method <sup>37</sup>. The basin falls under the hot 65 and sub-humid monsoonal climatic region. The average annual rainfall is 1444 mm. Maximum (82%) 66 67 rainfall occurs from June to September. There is a short rainy season in spring, March to May due to 68 NW disturbance. The estimated runoff of this basin area in monsoon time is 693.34 mm. which is also 69 a significant factor for controlling soil erosion potentiality.

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## Fig: 1 Study Area map

## 74 3. MATERIALS AND METHODS

For constructing morphoetric surface and its categorization, 18 associated parameters have been
taken into consideration. All the parameters have been described in table 1. All these data layers
have been developed from Toposheets (SOI), satellite imageries and SRTM DEM of United States
Gelogical Survey (USGS).

## 79 Table 1: Parameters and formulae of parameters

Parameters	Formulae
Bifurcation ratio(Rb)	Ratio of number of stream segments of one order to the number of the next higher order <sup>[38]</sup>
Stream frequency (Sf)	Total number of streams/Basin area [39]
Drainage density (Dd)	Total stream length/Basin area <sup>[40]</sup>
Drainage Texture (Dt)	Total number of stream in all segments /Basin perimeter <sup>[39]</sup>

Drainage Intensity (Di)	Drainage frequency/drainage density; Di=Sf/ Dd <sup>[40]</sup>
Infiltration Number (In)	In=Sf * Dd <sup>[40]</sup>
Areal symmetry (Aa)	Area of the low area of the stream/ Area of the more area of the stream <sup>40</sup>
Constant of Channel Maintenance(CCM)	1/Drainage density
Length of overland flow (Lof)	1/Dd*0.5 <sup>39</sup>
Form factor (Ff)	Basin area/Basin length <sup>2 [39]</sup>
Elevation (H)	Height at MSL.
Relief ratio(Rh)	Basin relief/Basin length <sup>[38]</sup>
Relative relief (Rhp)	Basin relief/(Perimeter × 100) <sup>[42]</sup>
Hypsometric Integral(Hi)	Ratio between Elevation and area
Average slope (s)	$\tan\theta = N \times \frac{I}{636.6}$
	N=number of contour cuttings per miles or km.,I=contour interval, 636.6=constant <sup>[37]</sup>
Ruggedness number (Rn)	Basin relief * Drainage density [43]
Dissection index (DI)	DI= Dd * (H /1000) <sup>[44]</sup>
Hydraulic Gradient (HG)	HG=(hc-hf)/D*100 <sup>(145)</sup>
-	hc=Height at source, hf=height at confluence
	D= Distance between source and confluence

#### **3.1 Methodology for Constructing Morphometric Models:**

82 For identifying morphoetric surface and its categorization, three models have been constructed, 83 namely relief diversity model (RDM), drainage diversity model (DDM), morphometric diversity model 84 (MDM). Separately these have been made for illustrating distinct characters of the drainage scape, 85 relief scape and as a whole Morphometric surface. While selecting parameters some of the 86 parameters are common because these can directly or indirectly represent both relief and drainage 87 diversity. For example, birfurcation ratio is basically a drainage parameter but high bifurcation ratio 88 indicates sudden change of the break of slope and thereby represents relief character also. Similarly, 89 Ruggedness number is basically a relief parameter but it is highly associated with terrain ruggedness 90 and can determine the arrangement of drainage network. Another motto behind such approach is to 91 understand the role of drainage and relief parameters for bringing morphometric diversity over the 92 basin area. Fractal analysis, discriminant analysis etc. were previously used by different scholars <sup>[46,47]</sup>for landscape identification, but present spatial approach can also be used for making such 93 94 differences over surface. Two methods are mainly used for developing all the models, first one is 95 simple Linear Combination (LC) method, where selected parameters are combined using raster calculator tool of Arc GIS 9.3 and another is Weighted Linear Combination (WLC) method of Carver 96 <sup>[48]</sup> and Eastman <sup>[49]</sup> where combination of the parameters are made by emphasizing their relative 97 importance. For calculating the weight of each parameter, PCA based correlation matrix is used and 98 99 data standardization for the same is done using equation 1 and 2. All the layers adopted here have 100 been converted into raster form because execution of compositing of layers is raster based. Two tier 101 weighting has been done for each layer. One is intra parameter ranking based on 10 point scale. 102 Selection of scale point should be based on level of accuracy required for work. Greater number of 103 classes will provide precision in final output. Here 10 classes have been done for describing the entire

104 spectrum of variable into manageable form. Maximum rank (e.g. 10) is provided to that class having 105 maximum diversity. Secondly, weight is assigned to each parameter based on their relative 106 importance to the aim. Weightage of each attribute has been defined objectively (see table 4 and5) 107 considering the degree of correlation of each driving factor with each other. The logic behind this 108 consideration is that highly correlated parameter maximally explains the spatial dissimilarity. 109 Normalization of respective weight (values of r for respective parameters) based on dimension index 110 has been done for considering it in a scientific scale. The result of each normalized value is called 111 attribute weight.

112 Expression of weight calculation is as follows:

113 
$$w_j = \frac{a_{j_r}}{\sum_{j=1}^n j_r}$$
 ..... (Eq. 1)

114 wj=weight of jth parameter; ajr= correlation coefficient of jth attribute;  $\Sigma$  jr = summation of correlation

of all jth variable.

116 Rank of all sub classes under each attribute is then multiplied by the defined weight of each individual

- 117 attribute. This function can be presented using the following formula.
- 118 WLC=  $\sum_{j=1}^{n} a_{ij} w_j$  ..... (Eq. 2)

Where, aij= ith rank of jth attribute; wj= weightage of jth attribute.

120 This weighted linear combination has been done using raster calculator tool in Arc GIS environment.

121 Table 2 shows selected spatial data layers of different models. Table 3 depicts the selected 10

122 parameters for relief diversity model i.e. Ruggedness number, dissection index, elevation, relative

123 relief, relief ratio, slope, hydraulic gradient, regional sinuosity, asymmetry factor and bifurcation ratio,

124 their subclasses, sub classes ranges, subclass ranking and weight of the parameters following the 10

- 125 points scale and PCA based WLC. For drainage and overall morphometric diversity model, same type
- 126 of process has been adopted but not described here separately. Table 4 shows logic behind 10 point
- 127 scale distribution among the intra classes and normalization process for final weight generation. For
- 128 example, slope is one of the parameters for this work. Here, 10 weights are assigned to highest
- 129 degree of slope class because it causes maximum diversity over surface. Table 5 and 6 describe the
- 130 same for drainage and morphometric diversity models.

#### 131 **Table 2** : Name and number of parameters used for different diversity models.

Type of model	Number of parameters	Name of the parameters
Relief diversity	10	Ruggedness number, dissection index, elevation, relative relief, relief ratio, slope, hydraulic gradient, regional sinuosity, asymmetry factor and bifurcation ratio
Drainage diversity	13	Drainage density, drainage frequency, Bifurcation ratio, relative relief, drainage intensity, drainage texture, regional sinuosity, junction frequency, constant channel maintenance, length of over land flow, infiltration number, form factor and ruggedness number.

Morphometric16Asymmetry, Stream Density, Stream Frequency, Drainage<br/>intensity, Dissection index, Drainage Texture, Form Factor,<br/>Hydraulic gradient (HG), Infiltration Number, Junction Frequency,<br/>Bifurcation Ratio, Relief Ratio, Relative Relief, Ruggedness<br/>Number, Elevation and Slope.

	of the parameters for construction relief	Ranks and weight of the	: Example of Assigning	135 Table 3
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136 diversity model

Parameters	Sub classes	Given rank	Weights of parameters	Parameters	Sub classes	Given rank	Weights of parameters
	0.095-	1	•		0.26-	1	
	0.16				0.57		
	0.16-	2			0.57-	2	
	0.20				0.82		
	0.20-	3			0.82-	3	
Dissection index	0.23		0.87		1.02		0.52
	0.23-	4		HG	1.02-	4	
	0.26	_			1.23	_	
	0.26-	5			1.23-	5	
	0.29	_			1.47	<u>^</u>	
	0.29-	6			1.47-	6	
	0.33	7			1.74 1.74-	7	
	0.33- 0.37	1			2.03	1	
	0.37	8			2.03	8	
	0.37-	0			2.03-	0	
	0.41-	9			2.30	9	
	0.47	3			2.74	3	
	0.47-	10			2.74-	10	
	0.58	10			3.32	10	
	0.00-	1			62-74	1	
	0.005					·	
	0.005-	2			74-83	2	
	0.011						
	0.11-	3			83-91	3	
Ruggedness	0.015		0.77	Elevation			
number	0.015-	4			91-99	4	0.58
	0.19						
	0.19-	5			99-106	5	
	0.024						
	0.024-	6			106-	6	
	0.029				113		
	0.029-	7			113-	7	
	0.034	0			121	0	
	0.034-	8			121-	8	
	0.041	9			129	9	
	0.041- 0.05	9			129- 138	9	
	0.05	10			138-	10	
	0.063	10			155	10	
	0.005	1			0-1.06	10	
	0-15	2			1.06-	9	
		-			1.11	~	

	15-21	3			1.11-	8	
Relative relief	21-26	4	1	Sinuosity	1.18 1.18-	7	0.32
					1.27		0.52
	26-30	5			1.27- 1.37	6	
	30-34	6			1.37-	5	
	34-39	7			1.48 1.48-	4	
					1.6		
	39-44	8			1.6- 1.72	3	
	44-48	9			1.72- 1.86	2	
	48-54	10			1.86- 1.99	1	
	0-0.38	1			0.29-	1	
	0.38-	2			0.35 0.35-	2	
	0.68				0.4		
Slope	0.68-0.1	3	0.79	Asymmetry Factor	0.4- 0.44	3	0.59
	0.1-1.39	4	0.75	1 40101	0.44-	4	0.00
	4.00	_			0.49	_	
	1.39- 1.79	5			0.49- 0.54	5	
	1.79- 2.12	6			0.54- 0.59	6	
	2.12-	7			0.59-	7	
	2.49 2.49-	8			0.63 0.63-	8	
	2.93 2.93-	9			0.68 0.68-	9	
	3.37				0.72		
	3.37- 4.05	10			0.72- 0.78	10	
	0.003- 0.011	1			1.94- 2.29	1	
	0.011-	2			2.29-	2	
Relief	0.022	3	0.64	Bifurcation ratio	2.73 2.73-	3	
ratio	0.038		0.0.1		3.18		a = 1
	0.038- 0.055	4			3.18- 3.54	4	0.71
	0.055-	5			3.54-	5	
	0.073	C			3.83	C	
	0.073- 0.089	6			3.83- 4.22	6	
	0.089-	7			4.22-	7	
	0.11 0.11-	8			4.69 4.69-	8	
	0.12				5.17		
	0.12- 0.14	9			5.17- 5.62	9	
	0.14- 0.15	10			5.62- 6.08	10	

- **Table 4**: Rank distribution, logic behind rank distribution, correlation score and weighted
- 140 score of Relief Diversity model.

Parameters	Rank	Logic	Total correlation score(Xi)	Weighted score Xi/Ximax
Dissection index	10 rank at highly dissected area	Extreme dissected area is a very good indicator of relief diversity	4.03	0.87
Ruggedness number	10 rank at maximum rugged topography	More value of Ruggedness number enhance the relief diversity	3.57	0.77
Relative relief	10 rank at highest relative relief	Highest value of relative relief increase the relief diversity	4.63	1
Slope	10 rank at steep slope	Greater value of slope promote the relief diversity	3.65	0.79
Relief Ratio	10 rank at maximum relief ratio	More value of relief ratio demarcated more relief diversity	2.95	0.64
HG	10 rank at highest value of Hydraulic gradient	Higher value of hydraulic gradient shows greater relief diversity	2.38	0.52
Elevation	10 rank at maximum relief	Greater value of relief indicate the more relief diversity	2.66	0.58
Sinuosity	10 rank at less sinuous region	Where sinuosity is low, relief diversity is high	1.49	0.32
Asymmetry Factor	10 rank at more value of asymmetry	More asymmetry shows more relief diversity	2.71	0.59
Bifurcation ratio	10 rank at highest value of Bifurcation ratio	Higher value of Bifurcation ratio denotes the greater relief diversity	3.30	0.71

# 143Table 5 : Rank distribution, logic behind rank distribution, correlation score and weighted144score of Drainage Diversity.

Parameters	Rank	Logic	Total correlation score(Xi)	Weighted score Xi/Ximax
Drainage frequency	10 rank at maximum drainage frequency	Where drainage frequency is maximum Drainage diversity also maximum	5.40	0.914
Drainage density	10 rank at highest drainage density	High drainage density indicates higher drainage diversity.	5.91	1
constant channel	10 rank at more CCM value	High CCM value favorable for more	4.27	0.72

maintenance( CCM)		Drainage diversity		
Drainge intensity	10 rank at greater intensity	Maximum Drainage Intensity lead to drainage diversity	3.04	0.52
Length of over land flow(Lof)	10 rank at maximum value of Lof	High Lof indicates higher drainage diversity	5.68	0.96
6.Drainagre texture(Dt)	10 rank at maximum drainage texture	high value of DT maximized drainage diversity	5.57	0.94
7.Junction frequency	10 rank at highest junction frequency	More junction frequency positively effect on drainage diversity	3.18	0.54
Regional sinuosity	10 rank at lowest sinuosity	Drainage diversity is greater at less sinuous river zone	1.60	0.27
Infiltration number(In)	10 rank at more IN	More In value indicates more drainage diversity	4.54	0.77
Bifurcation ratio(Rb)	10 rank at highest Bifurcation ratio	Frequently flood occurrence observed in high drainage diversity area.	3.40	0.58
Form factor(Ff)	10 rank at highest form factor	Higher value of form factor highlights the absolute drainage diversity	2.82	0.48
Relative relief	10 rank at highest relative relief	High elevation represent greater drainage diversity	2.97	0.50
Ruggedness number	10 rank at greater ruggedness number	Drainage diversity is high in rough topography	5.76	0.97

# 147Table 6: Rank distribution, logic behind rank distribution, correlation score and weighted148score of Morphometric Diversity.

Parameters	Rank	Logic	Total correlation score (Xi)	Weighted score (Xi/Xi max)
Asymmetry	10 rank at high asymmetric value	High asymmetric value increase morphometric diversity	3.89	0.62
Drainage Density	10 rank at highest stream density	High stream density indicates higher morphometric diversity.	5.28	0.84
Stream Frequency	10 rank at maximum stream frequency	Where stream frequency is maximum morphometric diversity also maximum	5.17	0.83
Drainage intensity	10 rank at greater intensity	greater Drainage Intensity lead to more diversity	2.80	0.45
Dissection index	10 rank at higher dissected value	More dissected zone indicates more morphometric diversity	5.06	0.81
Drainage Texture	10 rank at maximum drainage texture	high value of drainage texture maximized morphometric diversity	5.05	0.81

Form Factor	10 rank at highest form factor	Higher value of form factor highlights morphometric variety	3.84	0.61
Hydraulic gradient	10 rank at greater HG value	More HG value enhance the differentiation of morphometric diversity	3.07	0.49
Infiltration Number	10 rank at more IN value	More In value indicates more morphometric diversity	4.11	0.65
Junction Frequency	10 rank at highest junction frequency	More junction frequency positively effect on morphometric diversity	4.27	0.68
Bifurcation Ratio	10 rank at highest Bifurcation ratio	Frequently flood occurrence observed in high morphometric diversity area.	5.38	0.86
Relief Ratio	10 rank at highest relief ratio	High relief ratio represent greater morphometric diversity	4.05	0.65
Relative Relief	10 rank at highest relative relief	High relative relief indicates higher morphometric diversity	5.74	0.91
Ruggedness Number	10 rank at greater ruggedness number	morphometric diversity is high in rough topography	6.27	1
Slope	10 rank at steep slope	Steep slope positively control morphometric diversity	4.22	0.67
Elevation	10 rank at high elevation	Morphometric diversity influenced by high elevation	4.41	0.70



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Fig: 2 Work Flow Diagram

152 In figure 2 detail methodological flow for constructing relief diversity model has been made. For other

153 cases same approaches have been adopted as shown here but most of the parameters are different.

## 154 4. RESULT AND DISCUSSION

#### 155 4.1 Spatial data layers used for the models

A brief account of some morphometric characteristics of Kushkarani river basin are as follows: total length of the stream is 33.83 km., bifurcation ration is 4.027, length of over land low is 0.88, constant of channel maintenance is 0.568, sinuosity is 1.33, form factor of the basin is 0.268, drainage density is 1.76 km/sq km. drainage texture is 4.36, drainage intensity is 3.75, hypsometric integral (HI) is 0.48, relative relief is 93m., dissection index (DI) is 0.808, and ruggedness number is 0.062. All these values are representative to the entire basin but spatial variation is very strong.

162 Figure 3 (A-S) represent different spatial data layers used for different diversity model. Different data 163 layers behave differently towards the objective. For example in case of ruggedness number, relative 164 relief, drainage density, intensity, frequency these are high in the upper catchment of the basin 165 whereas it low in the same area i.e. form factor, bifurcation ratio and relief ratio. In the same way 166 differential behavior is also noticeable at the middle and lower catchment of the basin. For example 167 ruggedness number for this basin ranges from 0 to 0.062 (Fig. 3A). Few patches of upper and middle 168 catchments represent high range of value (0.05-0.06) covering an area of 11.6% to total area. 169 Dissection index(DI) ranges from 0.09 to 0.576 and high value is recognized at the upper catchment 170 and left portion of the middle catchment (Fig. 3B). Position of scattered hillock like landform has 171 generated this situation. Relative relief value ranges from 10-54m. (Fig. 3D) and it is high in those part 172 where dissection is high as mentioned earlier. Slope of this basin ranges from 0-4.05° and high 173 degree of slope is noticed at the upper fringe of the upper catchment (Fig. 3F). Figure 3H indicates 174 that sinuosity value ranges from 0-1.98 and no definite trend is identified over the basin as the entire 175 basin lies over the plateau fringe area. Bifurcation ratio (range: 1.93-6.07) is relatively higher in the 176 lower middle and upper part of the lower catchment (Fig. 3J). Since these areas, volume of water and 177 related functions may be energized. Drainage density of this basin ranges from 0-1.4 and it is found in 178 scattered patches over the basin (Fig. 3K).

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(A) Ruggedness Number









(E) Relief Ratio

(F) Slope





(G) Hydraulic Gradient

(H) Sinuosity Index



## 188 189

(I) Asymmetry Factor







Fig 3 (A –S) shows the spatial data layers for constructing the models. The parameters which are used for constructing the three models are described separately in table 2 -4.

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## 204 4.2 Relief diversity model (RDM)





## 206Fig 4 (a) Weighted Relief Diversity(b) Weighted Classified Relief Diversity

## 207 PCA based weights composite equation for Relief Diversity

208	<b>RDM=</b> (Rn*0.87)+(DI*0.77)+(Rhp*1)+(S*0.79)+(Rh*0.64)+(HG*0.52)+(H*0.58)+(Si*0.32)+(Aa*0.59)+(Rb*0.71)
209	(Eq.3)

210

### 211 **Table 7**: Area and percentage of area under different weighted Relief Diversity classes.

Relief diversity status	WLC score	Area extent (sq.km)	% of total area
Very low	50.81 - 67.97	24.88	14.47
Low	67.97 - 79.12	38.37	22.30
Moderate	79.13 - 90.34	46.65	27.12
High	90.34- 102.40	41.80	24.30
Very High	102.40- 125.68	20.30	11.81

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213 Relief diversity model mainly is constructed for identifying the unique geomorphological surface based 214 on relief characteristics in the river basin. Equation 3 represents the concise weighted linear 215 combination model of relief diversity. Figure 4(a) shows that very high relief diversity is found at the 216 upper catchment of the basin where ruggedness number, dissection index, elevation, relative relief 217 and slope are high. In this area, average slope is 2.02%, DI is 3.625, ruggedness number is 0.031. 218 Such conditions indicate that the landscape unit is characterized by greater soil erosion potentiality, 219 less chance of ground water recharge etc. Due to having greater potentiality of top soil loss and 220 therefore, agriculturally this region will not be potential enough. Tree feeling incidents will be more 221 frequent in this area. Out of the total area 11.81% area is found under very high relief diversity zone 222 (table 7). In case of very low relief diversity, it is found at the lower catchment of the basin where 223 ruggedness number, dissection index, elevation, relative relief, and slope are low. Out of the total 224 area 14.47% area are come under very low relief diversity zone. This region usually will be highly

- 225 potential for agricultural activities due to greater potentiality of soil moisture and less susceptibility to
- soil erosion. Surface and sub surface water potentiality will also support such economic activities.

## 227 **4.3 Drainage diversity model (DDM)**



#### 228

Fig 5 (a) Weighted Composite Drainage Diversity (b) Weighted Composite Classified Drainage
 Diversity

#### 231 PCA based weights composite equation for Drainage Diversity.

- 232  $DDM=(Sf^{*}0.91)+(Dd^{*}1)+(CCM^{*}0.72)+(Di^{*}0.52)+(Lof^{*}0.96)+(Dt^{*}0.94)+(Jf^{*}0.54)+(Si^{*}0.27)+(In^{*}0.77)+(Rb^{*}0.58)+(Ff^{*}0.48)+(Rhp^{*}0.50)+(Rn^{*}0.97)...(Eq. 4)$
- 233

#### 234 **Table 8**: Area and percentage of area under different Weighted Drainage Diversity classes.

Drainage Diversity	WLC score	Area extent (sq.km)	% of total area
Very low	13.02 - 26.40	14.85	8.63
Low	26.40 - 35.23	45.72	26.58
Moderate	35.23 - 42.50	50.18	29.18
High	42.50- 51.09	40.14	23.34
Very High	51.09 - 68.49	21.11	12.27

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236 Drainage diversity model has been prepared for identifying unique drainage landscape unit. Equation 237 4 depicts weighted linear composite model for drainage diversity model. Figure 5(a) and 5(b) 238 respectively illustrated the drainage diversity model in continuous and classified form. Very high 239 drainage diversity is observed at middle and upper catchment of the basin where drainage density (1-240 1.49 km/sq.km), drainage frequency (3.20-4.625 stream/sq.km), drainage texture (0.732-241 1.860/sq.km.), length of over land flow (0.340-0.713) and ruggedness number (0.048-0.0627) is high. 242 Table 8 shows that out of the total area of the basin 12.27% area falls under very high drainage diversity zone. This zone is characterized by large number of 1<sup>st</sup> and 2<sup>nd</sup> order streams and therefore 243 highly prone to soil loss. Constant erosion activities in this region are not also suitable for agriculture 244 245 activities and construction of settlement. It causes frequent felling of trees. Very low drainage diversity 246 is observed at the lower catchment of the basin where drainage diversity, drainage texture, drainage 247 frequency and ruggedness number are very low. Out of the total area 8.63% area (Table 8) comes 248 under very low drainage diversity zone. Dominance of rills and gullies accelerate the erosion process 249 and natural deforestation.

#### 250 **4.4** Morphometric Diversity models (MDM)



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Fig 6 (a) Weighted Morphometric Diversity

(b) Weighted classified Morphometric Diversity

#### 254 PCA based weights composite equation for Morphometric Diversity.

**MDM**=(Aa\*0.62)+(Dd\*0.84)+(Sf\*0.83)+(Di\*0.45)+(DI\*0.81)+(Dt\*0.81)+(Ff\*0.61)+(HG\*0.49)+(In\*0.68)+(Rb\*0 .86)+(Rh\*0.65)+(Rhp\*0.91)+(Rn\*1)+(S\*0.67)+(H\*0.70).....(Eq. 5)

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#### 258 **Table 9**: Area and percentage of area under different Morphometric diversity classes.

Morphometric Diversity	WLC score	Pixel count	Area extent (sq.km)	% of total area
1 (Lowest)	65.59-75.11	225	1.72222	1.00
2	75.11-84.64	724	5.54172	3.22
3	84.64-94.16	2593	19.84763	11.54
4	94.16-103.69	4206	32.19403	18.72
5	103.69-113.21	4983	38.14143	22.17
6	113.21-122.74	3880	29.69872	17.26
7	122.74-132.26	3220	24.64688	14.33
8	132.26-141.78	1509	11.55035	6.71
9	141.78-151.31	858	6.567398	3.82
10 (Highest)	151.31-160.84	273	2.089627	1.21
Total		22471	172	100

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260 Incorporating all the Morphometric parameters finally diversity in Morphometric surface has been 261 constructed (Eq. 5). Figure 6 (a and b) show continuous WLC and classified WLC model for 262 morphometric diversity model respectively. Classified model highlights ten distinct morphometric 263 surfaces with distinct morphometric characters (Fig. 6b). Table 9 represents the area under different 264 morphometric diversity classes, where 1.21% area comes under very high and 1.00% area comes 265 under very low diversity surface. Highly diversified morphometric surface is geomorphological 266 picturesque but accelerated rill and gully action is majorly responsible for high rate soil erosion, fast rate tree felling etc [36]. Soil erosion rate of this area is 6-18 tons/ha/year as per the estimation of Jha 267 268 and Kapat <sup>[50]</sup> .Least diversified surface mainly comes under relatively depressed surface at the 269 confluence region where deposition action is more active triggered by back thrust of flow due to rise of water level in Tilpara barrage <sup>[51]</sup>. This surface is submerged seasonally during monsoon period. Table 270

271 10 briefly describes the statistics of different morphometric parameters in different morphometric 272 diversity zones. From this table one can assume the differences of different morphometric surface 273 units. For example, at diversity zone 10 where diversity level is high and it is explained by a number of 274 parameters like high drainage density (0.9-1.40km./sq.km.), high dissection index (0.29-0.50/sq.km.), 275 high stream frequency (2.1-4.5/ sq.km.) etc. This part experiences late youth to mature stage of the 276 cycle of erosion as indicated by hypsometric integral (0.42-0.55). On the other hand lowly diversified 277 area is characterized by low drainage density (0-0.32km./sq.km.), low dissection index (0.09-278 0.1/sq.km.), low stream frequency (0-0.5/sq.km.) etc. This zone is located at the late mature to old

279 stage of the cycle of erosion (HI=0.35-0.45).

### 280 **Table 10: Parameter characters in different diversity clusters**

Mor pho metr ic Dive rsity	Dd( <mark>km/</mark> sq.k m)	Rh p ( <mark>in</mark> m)	Rn	DI	C C M	Si	Di	In	D t	L of	S( <mark>i</mark> n de gr ee )	S f( st re a m o. /s q. k m	H G	Hi	Ff	R b	Rh	H (i n m )
1	0- 0.32	10- 13	0- 00 6	0.0 9- 0.1	0- 6	1- 1. 98	0. 0 5	0- 0. 5	0- 0. 5	0- 0. 2	0	0- 0. 5	1. 5- 2	0. 35 - 0. 45	0. 1- 0. 52	2- 3. 5	0- 0.0 01	62 - 65
2	0.05 - 0.46	12- 16	0.0 02- 0.0 03	0.0 9- 0.3	0- 5. 5	1. 4- 1. 9	0. 3- 1	0- 2. 5	0. 5- 0. 9	0. 1- 0. 2	0- 0. 1	0. 5- 1. 2	1. 6- 2. 9	0. 39 - 0. 42	0. 75 - 0. 89	2- 4. 2	0.0 01- 0.0 5	63 - 69
3	0.2- 1	13- 20	0.0 02- 0.0 04	0.1 - 0.4 5	3- 1 3	1- 1. 9	0. 8- 3. 5	1- 5. 5	0. 5- 1. 1	0. 1- 0. 3	0. 2- 1	0. 2- 3. 5	1. 2- 2. 1	0. 36 - 0. 46	0. 12 - 0. 56	1. 9- 5	0.0 2- 0.0 5	63 - 75
4	0.1- 1.2	12- 22	0.0 03- 0.0 06	0.0 2- 0.4 2	0- 7	0. 4- 1. 5	0. 1- 4. 5	1. 5- 5	0. 1- 1. 5	0- 0. 3	0. 5- 1. 1	0. 3- 4	0. 26 - 1. 56	0. 36 - 0. 45	0. 12 - 0. 61	2. 3- 5. 2	0.0 02- 0.1 5	75 - 83
5	0.6- 1.3	15- 25	0.0 05- 0.0 06	0.0 9- 0.3	4- 1 3. 5	0. 8- 1. 6	0. 5- 6. 5	2. 5- 6. 5	0. 4- 1. 6	0. 1- 0. 6	0. 2- 1. 3	1- 3	0. 28 - 3. 1	0. 38 - 0. 46	0. 15 - 0. 54	2. 5- 5. 5	0.0 2- 0.0 6	80 - 10 0
6	0.7- 1.35	20- 45	0.0 07- 0.0 6	0.2 - 0.4	2- 8	0- 0. 8	2- 6	1. 5- 4. 5	0. 6- 1. 4	0. 2- 0. 6	0. 5- 2. 5	0. 1- 4. 5	2- 3	0. 4- 0. 5	0. 45 - 0. 9	3- 5. 5	0.0 6- 0.0 7	78 - 11 5
7	0.8- 1.39	35- 54	0.0 1-	0.3 -	3- 8.	0. 1-	2. 5-	3. 5-	0. 7-	0. 5-	2. 5-	0. 2-	1. 5-	0. 3-	0. 2-	2. 5-	0.0 5-	10 0-

			0.0 6	0.5 6	5	1. 6	7. 1	7. 4	1. 7	0. 7 1	3. 5	3. 5	3	0. 51	0. 9	6	0.6	12 1
8	0.65 - 1.36	40- 50	0.0 2- 0.0 4	0.3 5- 0.4 6	3- 1 0	0. 2- 1. 1	4- 6	4. 5- 7. 2	0. 7- 1. 6	0. 4- 0. 6 8	3- 4	2- 4. 2	1. 0- 3. 3	0. 35 - 0. 49	0. 45 -1	3. 5- 5. 1	0.0 4- 0.0 7	11 5- 13 2
9	0.7- 1.39	30- 53	0.0 3- 0.5 2	0.3 6- 0.5 1	2- 7	0. 1- 0. 9	5- 7. 1	3- 7. 1	1. 2- 1. 8	0. 5- 0. 7	3. 5- 4. 1	2. 1- 4. 5	2. 9- 3. 33	0. 45 - 0. 51	0. 65 - 1. 0	5. 1- 5 6	0.0 6- 0.0 8	12 0- 14 0
10	0.9- 1.40	25- 54	0.0 2- 0.0 56	0.2 9- 0.5 0	2. 5- 8	0. 15 - 0. 5	4. 5- 7	3. 5- 7. 4	1. 3- 1. 7	0. 5- 0. 7 1	3. 6- 4. 1	2. 6- 4. 6	3. 0- 3. 3	0. 42 - 0. 55	0. 63 - 1. 1	5- 5. 8	0.0 6- 0.0 8	13 5- 15 5

#### 282 **4.5** Spatial Association among the RDM, DDM and MDM

283 Correlation matrix has been carried out among three above mentioned diversity models to know the 284 role of drainage and relief parameters for making the surface diversified. Table 11 shows relief 285 diversity and morphometric diversity are strongly correlated as the correlation value is 0.94 and 286 drainage diversity is positively correlated with RDM and MDM as the correlation values are 0.445 and 287 0.207 respectively. Former two cases represent significant correlation at 0.01 level of significance. In 288 an ideal river basin, relief and drainage is highly associated, elevation, slope control flow form, flow 289 density etc. but at this present case association between relief and drainage is not so strong. Relation 290 is positive (r=o.207) and also significant at 0.1 level of significance between two but not statistically 291 acceptable. Low relief diversity over this plateau fringe river is the major reason behind such low 292 spatial variability. From this correlation structure, it is clear that for bringing morphometric diversity, 293 relief parameters play stronger role than drainage diversity and drainage diversity of the landscape is 294 also controlled by relief diversity.

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#### 296 **Table 11**: Correlation among three models

	Relief diversity	Drainage diversity	Morphometric diversity
Relief diversity	1	0.207	0.940
Drainage diversity		1	0.445
Morphometric diversity			1

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## 298 <mark>5. CONCLUSION</mark>

From the analysis of three models different unique drainage, relief and morphometric surface(s) have been recognized. Composite weighted score of the relevant parameters have been classified into equal 10 classes representing 10 morphometric surfaces. Range of value with greater magnitude of values is considered as most diversified surface with relatively dynamic processes and forms and vice versa. Each unit can be treated as unique in terms of terrain processes and forms. These are also

- 304 distinct in connection with potentiality for developing its own characteristics assemblages cultural
- 305 landscapes. Distinguished terrain processes, soil catena, soil loss, water retaining capacity, recharge,
- 306 surface water availability and runoff characters characterizes the economic activities of the study area
- 307 specially agricultural activities. So, not only for categorizing morphological units but also for devising
- 308 economic landscaping such units are important.

#### 309 **References**

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- Blankson EJ, Green BH. Use of landscape classification as an essential prerequisite to
   landscape evaluation. Landscape Urban Planning.1991;21:149-162.
- Otero PI et al. Landscape evaluation: Comparison of evaluation methods in a region of Spain.
   Journal of Environmental Management. 2007. In Press.
- Staler A N. Quantitative geomorphology of basins and channel networks. In: Chow, V.T.
   (Ed.), Handbook of applied hydrology. Mcgraw Hill Book Company, New York.1964.
- Mages NS, Jitheslal KV, Chandrasekar N. Geographical information system-based morphometric analysis of Bharathapuzha river basin, Kerala, India, Appl Water Sci .
   2013;3:467–477.
- Rastogi RA, Sharma TC. Quantitative analysis of drainage basin characteristics. Jour Soil and water Conservation in India. 1976; 26:18-25.
- Baker VR. Introduction: regional landforms analysis. In: M Short, RW Blair, Jr. (eds),
   Geomorphologyfrom space, a global overview of regional landforms. National Aeronautics
   and Space Administration, SP 486. US Government Printing Office, Washington DC. 1986:1–
   26.
  - Nautiyal MD. Morphometric analysis of a drainage basin, district Dehradun, Uttar Pradesh, J Indian Soc Remote Sens.1994; 22(4) :251–261.
- Nag SK, Chakraborty S. Influence of rock types and structures in the development of drainage network in hard rock area, J Indian Soc Remote Sens. 2003;31(1): 25–35.
- Rudraiah M, Govindaiah S, Srinivas VS. Morphometry using remote sensing and GIS techniques in the sub-basins of Kagna river basin, Gulburga district, Karnataka, India. J
   Indian Soc Remote Sens.2008; 36: 351–360.
- Magesh NS, Jitheshlal KV, Chandrasekar N, Jini KV. GIS based morphometric evaluation of
  Chimmini and Mupily watersheds, parts of Western Ghats, Thrissur District, Kerala, India.
  Earth Sci Inform. 2012; 5(2):111–121.
- 11. John Wilson JS, Chandrasekar N, Magesh NS. Morphometric analysis of major sub watersheds in Aiyar and Karai Pottanar Basin, Central Tamil Nadu, India using remote
   sensing and GIS techniques. Bonfring Int J Indus Eng Manag Sci. 2012; 2(1):8–15.
- 12. Van RA, Verstraeten G, Van OK, Govers G, Poesen J. Modeling mean annual sediment yield
   using a distributed approach. Earth Surf Process Landforms. 2001; 26(11):1221–1236.
  - 13. Fernandez C, Wu JQ, McCool DK, Stockle CO. Estimating water erosion and sediment yield with GIS, RUSLE, and SEDD. J Soil Water Conserv. 2003;58(3):128–136.
  - 14. Ganapuram Sreedhara et al. Mapping of groundwater potential zones in the Musi basin using remote sensing data and GIS. Advances in Engineering Software. 2009; 40: 506-518.
  - 15. Vasanthavigar M et al. Groundwater potential zoning in Thirumanimuttar sub-basin Tamilnadu, India a GIS and remote sensing approach. Geo-spatial Information Science .2011;14(1):17 26. DOI 10.1007/s11806-011-0422-2
- 347
   348
   16. Awawdeh M et al. Integrated GIS and remote sensing for mapping groundwater potentiality in 348 the Tulul al Ashaqif, Northeast Jordan. Arabian Journal of Geoscience. 2014;7(6):2377-2392.
- 349 17. Jasrotia AS, Kumar A, Singh R. Integrated remote sensing and GIS approach for delineation
   350 of groundwater potential zones using aquifer parameters in Devak and Rui watershed of
   351 Jammu and Kashmir, India. Arabian Journal of Geosciences.2016; 9: 304.

352 18. Ibrahim-Bathis K, Ahmed S. Geospatial technology for delineating groundwater potential 353 zones in Doddahalla watershed of Chitradurga district, India. The Egyptian Journal of Remote 354 Sensing and Space Science.2016; 19(2): 223-234. 355 19. Gajbhiye S., Mishra S. K., Pandey A. Prioritizing erosion-prone area through morphometric 356 analysis: an RS and GIS perspective .Appl Water Sci. 2014:4:51-61.DOI 357 10.1007/s13201-013-0129-7 358 20. Gopal KG, Saha S. Identification of soil erosion susceptible areas in Hinglo river basin, 359 Eastern India based on geo-statistics. Universal Journal of Environmental Research and 360 Technology. 2015;5(3):152-164. 361 21. Sarkar BC. A GIS approach to morphometric analysis of Damodar river basin and 362 groundwater potentiality mapping in Jharia coalfield. 2011;26 (1-4):1-70. 363 22. Avinash K, Deepika B, Jayappa K S. Basin geomorphology and drainage morphometry 364 parameters used as indicators for groundwater prospect: Insight from geographical information system (GIS) technique. Journal of Earth Science. 214:25(6): 1018-1032. 365 366 doi:10.1007/s12583-0140505-8 367 23. Nidhi K , Chowdary VM, Tiwari KN, Shinde V, Dadhwal VK. Assessment of surface water potential using morphometry and curve number-based approaches. 368 Geocarto 369 International.2016. www.tandfonline.com/doi/pdf/10.1080/10106049.2016.1195889 370 24. Dahiphale P, Singh PK, Yadav KK. Morphometric analysis of sub-basins in jaisamand 371 catchment using geographical information system. International Journal of Research in 372 Engineering & Technology.2014; 2 (6) : 189-202. 373 25. Rai PK., Mohan K., Mishra S, Ahmad A, Mishra VN. A GIS-based approach in drainage 374 morphometric analysis of Kanhar River Basin, India. Appl Water Sci.2014: 1-16. DOI 375 10.1007/s13201-014-0238-y. 26. Khatun S and Pal S. Identification of prospective surface water available zones with multi 376 377 criteria decision approach in Kushkarani river basin of Eastern India, Archives of Current 378 Research International. 2016;4(4): 1-20. DOI:10.9734/ACRI/2016/27651 379 27. Sures M. et al. Prioritization of watersheds using morphometric parameters and assessment 380 of Surface Water Potential using Remote Sensing. Journal of the Indian Society of Remote 381 Sensing.2004;32(3):249-259. 382 28. Thakkar A K., Dhiman S. D. Morphometric analysis and prioritization of mini watersheds in 383 Mohr watershed, Gujarat using remote sensing and GIS techniques. Journal of the Indian Society of Remote Sensing.2007; 35(4):321-329.DOI: 10.1007/BF02990787 384 385 29. Amani M, Safaviyan A. Sub-basins prioritization using morphometric analysis-remote sensing 386 technique and GIS-Golestan-Iran. 2015;38:56-65. International Letters of Natural Sciences 387 Online: 2015:05-06. 388 30. Rahaman SA, Ajeez SA, Aruchamy S, Jegankumar R. Prioritization of sub watershed based 389 on morphometric characteristics using fuzzy analytical hierarchy process and geographical 390 information system - A study of Kallar watershed, Tamil Nadu. Aquatic Procedia. 391 2015;4:1322-1330. 392 31. Clark JJ and Wilcock PR. Effects of land use change on channel morphology in Northeastern 393 Puerto Rico. Bulletin, Geol. Society of America. 2000;112(12):1763-1777. 394 32. Godoy, Pinto MD, Lacerda de, Drude Luiz .River-island morphological response to basin 395 land-use change within the Jaguaribe river estuary, NE Brazil. Journal of Coastal 396 Research.2014;30(2):39. 397 33. Khatun S and Pal S. Analysis of regional hypsometric integral to identify landscape evolution 398 in Kushkarani river basin. Journal of Geography, Environment and Earth Science 399 International. 2016; 6(3): 1-17. 400 401 34. Bagchi K, Mukerjee KN. Diagnostic survey of West Bengal(s), dept. of geography. Calcutta 402 University, Pantg Delta & Rarh Bengal. 1983;17(19):42-58.

403	35. Chakrabarty SC. Some consideration on the evolution of physiogrphy of Bengal, in
404	Chattopadhyay B. (ed.), West Bengal, Geog. Inst., Presidency College, Calcutta.1970;20.
405	36. Pal S. Identification of Soil Erosion Vulnerable Areas in Chandrabhaga River Basin: a Multi
406	criteria Decision Approach, Model. Earth Syst. Environ. 2016; 2(5):1-11. DOI
407	10.1007/s40808015 0052-z
408	37. Wentworth CK. A simplified method of determining the average slope of land surfaces.
409	American Journal of Science. 1930;21:184-194.
410	38. Schumn SA. Evolution of drainage systems and slopes in badlands at Perth Amboy, New
411	Jersey. Geol. Soc. Am. Bull.1956;67:597-646.
412	39. Horton RE. Erosional development of streams and their drainage basins, hydrological
413	approach to quantitative morphology. Bull. Geophys. Soc. Am. 1945; 56:275-370.
414	40. Faniran A. The index of drainage intensity - A provisional new drainage factor. Australian
415	Journal of Science. 1968; 31: 328-330.
416	41. Cox RT. Analysis of drainage-basin symmetry as a rapid technique to identify areas of
417	possible quaternary tilt-block tectonics: an example from the Mississippi Embayment.
418	Geological Society of America Bulletin. 1994; 106: 571- 581.
419	42. Sarangi A, Madramootoo C, Enright MP. Development of user interface in Arc GIS for
420	estimation of watershed geomorphology. The Canadian society for engineering in
421	agricultural. Food and Biological Sytsems. 2003:03-120.
422	43. Miller JR, Ritter DF, Kochel RC. Morphometric assessment of lithologic controls on drainage
423	basin evolution in the Crawford Upland, South-Central Indiana. American Journal of
424	Science.1990;290:569-599. http://dx.doi.org/10.2475/ajs.290.5.569
425	44. Patton PC and Baker VR. Morphometry and floods in small drainage basins.1976;12(5): 941-
426	952
427	45. Strahler AN. Quantitative analysis of watershed geomorphology. Transactions, American
428	Geophysical Union. 1957;38 (6):913-920.
429	46. Wu et al. Gap shape classification using landscape indices and multivariate statistics.
430	Scientific Report 6. 2016.doi:10.1038/srep38217
431	47. Klinger R, Schwanghart W, Schutt B. Landscape classification using principal component
432	analysis and fuzzy classification: Archaeological sites and their natural surroundings in
433	Central Mongolia, Geoarchaeology. 2011; 3:213-233
434	48. Carver SJ. Integrating multi-criteria evaluation with geographical information systems.
435	International Journal of Geographical Information Systems. 1991; 5(3):321–339.
436	49. Eastman JR. Idrisi for windows, version 2.0: Tutorial exercises, graduate school of
437	geography—Clark University, Worcester, MA; 1997.
438	50. Jha VC, Kapat S. Gully erosion and its implications on land use, a case study. In JHA, V.C.
439	(Ed). Land degradation and desertification. Publ., Jaipur and New Delhi. 2003;156-178.
440	51. Pal S. Identification of soil erosion vulnerable areas in Chandrabhaga river basin: a multi-
441	criteria decision approach, model. Earth Syst. Environ. 2016; 2(5):1-11; DOI 10.1007/s40808-
442	015-0052-z