Original Research Article

Categorization of Morphometric Surface through Morphometric Diversity Analysis in Kushkarani River Basin of Eastern India

5

3

4

1 2

6 ABSTRACT

7 The present work attempts to build up and categorize morphometric surface of distinct identities and 8 potentialities considering 16 morphometric parameters as spatial data layers in GIS environment. For 9 constructing the morphometric diversity model relative importance of the selected parameters is 10 assessed and weighted compositing is being made. Separately, relief and drainage diversity models 11 have been prepared and finally entire surface is categorized into some distinct morphometric sub 12 units having their own terrain characteristics and associated cultural features. Out of total river basin 13 only 11% area comes under highly diversified surface and maximum part of the area comes under 14 quite monotonous surface with least variation. Highly diversified surface is prone to rilling, gullying and 15 consequently soil erosion. It is also tried to identify whether relief and drainage model exert significant 16 impact on morphometric diversity model. It is identified that relief diversity model carries significant 17 impact on morphometric diversity model as indicated by high correlation coefficient (r=0.94).

Keywords: Mophometric surface, Weighting of parameters, Weighted linear combination (WLC),
 Drainage diversity model (DDM), Relief diversity model (RDM) and Morphometric diversity model
 (MDM).

21 **1. INTRODUCTION**

Landscape is an essentially visual phenomenon or as a particular constitution of topography, land 22 use, vegetation cover and settlement pattern ^{1,2} and morphometric surface is one of the fundamental 23 terrain base on which cultural and ecological conditions can be draped. Morphometric analysis 24 25 includes a good number of quantitative indicators based on which terrain units can be recognized and 26 each such terrain units (unit) provides a strong base of (for) analyzing terrain processes and associated cultural fabrics ^{3,4}. Previous studies focused on morphometric characteristics of various 27 drainage basins using GIS and remote sensing technique for terrain analysis ⁵⁻¹¹. The Morphometric 28 29 parameter (parameters) calculated in the study are flow directions, flow accumulation, upstream and downstream, stream link, stream network, stream order, and digital elevation models (DEMs) etc.^{12,13}. 30 31 Generally in this studies emphasize segregated statistical analysis or mapping of the individual morphometric parameters. Among then some studies correlate morphometric parameters and 32 potential ground water investigation ¹⁴⁻²⁰, recharge estimation²¹⁻²³ and surface water availability²⁴⁻²⁶ 33 etc. Some works (studies) laid importance on sub basin priority for promising planning unit ²⁷⁻³⁰. The 34 35 present work pursues how to indentify morphometric surface units using different morphometric 36 parameters as spatial data layers. Primarily, drainage and relief morphometric surface(s) are tried to

37 (delete) recognized and finally morphometric surface is also being identified. Here attempt is also
 38 being taken (made) regarding prioritizing incorporated parameters considering their importance
 39 towards bringing morphometric diversity.

40 Morphometric diversity analysis actually helps to find out some nearly homogenous landscape units 41 which are uniform in morphometric characteristics. Each unit can be considered as a homogenous 42 hydro-geomorphic spatial unit with unique nature and potential viability of land use. Landscape unit 43 with highly diversified morphometric characteristics is healthy in geomorphic processes and forms and 44 active in term(s) of ongoing terrain processes. On the contrary, lowly diversified landscape unit is 45 characterized by relatively monotonous processes and forms. Each landscape unit is uniquely 46 potential for resources and economic activities. This fact becomes highly vivid when range of 47 morphometric differences is very high and size scale is very large. In case of small basin over 48 homogenous physiographic unit, it is not so clear and therefore, not so effective for dictating the LULC characteristics ^{31, 32}. But some indirect influences can be streamlined. For example, morphometric 49 50 diversity unit has its potentiality to characterize soil loss processes, runoff processes, recharge 51 activities etc. These are often responsible for controlling LULC and livelihood opportunities. 52 Hypsometric surface is well explained in relation to geology of the present study area using hypsometric integral (HI) of different sub basins by Khatun and Pal ³³. This surface 53 54 envisages the erosion processes, soil loss, sediment deposition pattern etc.

55

56 **2. STUDY AREA**

57 Kushkarani river (length: 35 km) basin, covering an 172 sq km area (see fig-1), is a sub basin of 58 Mayurakshi river system located mainly over the western part of Chottonagpur plateau fringe in 59 Birbhum district of West Bengal and Jamtara district of Jharkhand with 2354' 36" N to 2400' N 60 latitudes and 87°14'24" E to 87°30' E longitudes. The total basin area comes under rarh tract topography ³⁴ with lateritic soil formation ³⁵ which is mainly received by flowing rivers from 61 62 Chottonagpur plateau. This river pours into Tilpara barrage which is located over master stream Mayurakshi river ³⁶. The elevation of this catchment varies from 155 m. (at the source region) to 62 63 64 m. (at the confluence region). Maximum area of the basin shows rugged topography with an average 65 elevation of 108 m. Average slope of the basin is 1-4 degree whereas it is <1 degree in the confluence part of the basin measured as per Wentworth's method ³⁷. The basin falls under the hot 66 67 and sub-humid monsoonal climatic region. The average annual rainfall is 1444 mm. Maximum (82%) 68 rainfall occurs from June to September. There is a short rainy season in spring, March to May due 69 to nor'wester (NW) disturbance. The estimated runoff of this basin area in monsoon time is 693.34 70 mm. which is also a significant factor for controlling soil erosion potentiality.

- 71
- 72



74

Fig: 1 Study Area map

75 3. MATERIALS AND METHODS

For constructing morphoetric surface and its categorization, 18 associated parameters have been taken into consideration. All the parameters have been described in table 1. All these data layers have been developed from Toposheets (SOI), satellite imageries and SRTM DEM of United States Gelogical Survey (USGS).

80 Table 1: Parameters and formulae of parameters

Parameters	Formulae
Bifurcation ratio(Rb)	Ratio of number of stream segments of one order to the number of the next higher order ^[38]
Stream frequency (Sf)	Total number of streams/Basin area [39]
Drainage density (Dd)	Total stream length/Basin area ^[40]
Drainage Texture (Dt)	Total number of stream in all segments /Basin perimeter ^[39]

Drainage Intensity (Di)	Drainage frequency/drainage density; Di=Sf/ Dd ^[40]
Infiltration Number (In)	In=Sf * Dd ^[40]
Areal symmetry (Aa)	Area of the low area of the stream/ Area of the more area of the stream $^{\rm 40}$
Constant of Channel Maintenance(CCM)	1/Drainage density
Length of overland flow (Lof)	1/Dd*0.5 ³⁹
Form factor (Ff)	Basin area/Basin length ^{2 [39]}
Elevation (H)	Height at MSL.
Relief ratio(Rh)	Basin relief/Basin length ^[38]
Relative relief (Rhp)	Basin relief/(Perimeter × 100) ^[42]
Hypsometric Integral(Hi)	Ratio between Elevation and area
Average slope (s)	$\tan\theta = N \times \frac{I}{636.6}$
	N=number of contour cuttings per miles or km.,I=contour interval, 636.6=constant ^[37]
Ruggedness number (Rn)	Basin relief * Drainage density ^[43]
Dissection index (DI)	DI= Dd * (H /1000) ^[44]
Hydraulic Gradient (HG)	HG=(hc-hf)/D*100 ^[45]
	hc=Height at source, hf=height at confluence
	D= Distance between source and confluence

82 **3.1 Methodology for Constructing Morphometric Models:**

83 For identifying morphoetric surface and its categorization, three models have been constructed, 84 namely relief diversity model (RDM), drainage diversity model (DDM), morphometric diversity model 85 (MDM). Separately these have been made for illustrating distinct characters of the drainage scape, 86 relief scape and as a whole Morphometric surface. While selecting parameters some of the 87 parameters are common because these can directly or indirectly represent both relief and drainage 88 diversity. For example, birfurcation ratio is basically a drainage parameter but high bifurcation ratio 89 indicates sudden change of the break of slope and thereby represents relief character also. Similarly, 90 Ruggedness number is basically a relief parameter but it is highly associated with terrain ruggedness 91 and can determine the arrangement of drainage network. Another motto behind such approach is to 92 understand the role of drainage and relief parameters for bringing morphometric diversity over the 93 basin area. Fractal analysis, discriminant analysis etc. were previously used by different scholars ^[46,47]for landscape identification, but present spatial approach can also be used for making such 94 differences over surface. Two methods are mainly used for developing all the models, first one is 95 96 simple Linear Combination (LC) method, where selected parameters are combined using raster calculator tool of Arc GIS 9.3 and another is Weighted Linear Combination (WLC) method of Carver 97 ^[48] and Eastman ^[49] where combination of the parameters are made by emphasizing their relative 98 99 importance. For calculating the weight of each parameter, PCA based correlation matrix is used and 100 data standardization for the same is done using equation 1 and 2. All the layers adopted here have 101 been converted into raster form because execution of compositing of layers is raster based. Two tier 102 weighting has been done for each layer. One is intra parameter ranking based on 10 point scale. 103 Selection of scale point should be based on level of accuracy required for work. Greater number of 104 classes will provide precision in final output. Here 10 classes have been done for describing the entire

105 spectrum of variable into manageable form. Maximum rank (e.g. 10) is provided to that class having 106 maximum diversity. Secondly, weight is assigned to each parameter based on their relative 107 importance to the aim. Weightage of each attribute has been defined objectively (see table 4 and5) 108 considering the degree of correlation of each driving factor with each other. The logic behind this 109 consideration is that highly correlated parameter maximally explains the spatial dissimilarity. 110 Normalization of respective weight (values of r for respective parameters) based on dimension index 111 has been done for considering it in a scientific scale. The result of each normalized value is called 112 attribute weight.

113 Expression of weight calculation is as follows:

114
$$w_j = \frac{a_{j_r}}{\sum_{j=1}^n j_r}$$
(Eq. 1)

115 wj=weight of jth parameter; ajr= correlation coefficient of jth attribute; Σ jr = summation of correlation

116 of all jth variable.

117 Rank of all sub classes under each attribute is then multiplied by the defined weight of each individual

- attribute. This function can be presented using the following formula.
- 119 WLC= $\sum_{j=1}^{n} a_{ij} w_j$ (Eq. 2)

Where, aij= ith rank of jth attribute; wj= weightage of jth attribute.

121 This weighted linear combination has been done using raster calculator tool in Arc GIS environment.

122 Table 2 shows selected spatial data layers of different models. Table 3 depicts the selected 10

123 parameters for relief diversity model i.e. Ruggedness number, dissection index, elevation, relative

124 relief, relief ratio, slope, hydraulic gradient, regional sinuosity, asymmetry factor and bifurcation ratio,

125 their subclasses, sub classes ranges, subclass ranking and weight of the parameters following the 10

- 126 points scale and PCA based WLC. For drainage and overall morphometric diversity model, same type
- 127 of process has been adopted but not described here separately. Table 4 shows logic behind 10 point
- scale distribution among the intra classes and normalization process for final weight generation. For
- 129 example, slope is one of the parameters for this work. Here, 10 weights are assigned to highest
- degree of slope class because it causes maximum diversity over surface. Table 5 and 6 describe the
- 131 same for drainage and morphometric diversity models.

132 **Table 2** : Name and number of parameters used for different diversity models.

Type of model	Number of parameters	Name of the parameters
Relief diversity	10	Ruggedness number, dissection index, elevation, relative relief, relief ratio, slope, hydraulic gradient, regional sinuosity, asymmetry factor and bifurcation ratio
Drainage diversity	13	Drainage density, drainage frequency, Bifurcation ratio, relative relief, drainage intensity, drainage texture, regional sinuosity, junction frequency, constant channel maintenance, length of over land flow, infiltration number, form factor and ruggedness number.

Morphometric	16	Asymmetry, Stream Density, Stream Frequency, Drainage
Diversity		intensity, Dissection index, Drainage Texture, Form Factor,
		Hydraulic gradient (HG), Infiltration Number, Junction Frequency,
		Bifurcation Ratio, Relief Ratio, Relative Relief, Ruggedness
		Number, Elevation and Slope.

	136	Table 3: Example	of Assigning Ranks	s and weight of the par	rameters for construction relie
--	-----	------------------	--------------------	-------------------------	---------------------------------

137 diversity model

Parameters	Sub classes	Given rank	Weights of parameters	Parameters	Sub classes	Given rank	Weights of parameters
	0.095-	1			0.26-	1	
	0.16				0.57		
	0.16- 0.20	2			0.57- 0.82	2	
	0.20-	3			0.82-	3	
Dissection index	0.23		0.87		1.02		0.52
	0.23-	4		HG	1.02-	4	
	0.26				1.23		
	0.26-	5			1.23-	5	
	0.29	_			1.47	<u>_</u>	
	0.29-	6			1.47-	6	
	0.33	7			1.74	7	
	0.33-	1			2.03	1	
	0.37-	8			2.03	8	
	0.41	Ū			2.30	0	
	0.41-	9			2.30-	9	
	0.47				2.74		
	0.47-	10			2.74-	10	
	0.58				3.32		
	0.00-	1			62-74	1	
	0.005	0			74.00	0	
	0.005- 0.011	2			74-83	2	
	0.11-	3			83-91	3	
Ruggedness	0.015		0.77	Elevation			
number	0.015-	4			91-99	4	0.58
	0.19	_				_	
	0.19- 0.024	5			99-106	5	
	0.024-	6			106-	6	
	0.029				113		
	0.029-	7			113-	7	
	0.034	0			121	0	
	0.034-	8			121-	8	
	0.041-	9			129	9	
	0.05	Ū			138	J	
	0.05-	10			138-	10	
	0.063				155		
	0	1			0-1.06	10	
	0-15	2			1.06- 1.11	9	

	4 - 64	•				•	
Relative relief	15-21	3	1	Sinuosity	1.11- 1.18	8	
	21-26	4			1.18- 1.27	7	0.32
	26-30	5			1.27-	6	
	30-34	6			1.37-	5	
	34-39	7			1.48	4	
	39-44	8			1.6 1.6-	3	
	11-18	9			1.72	2	
	40.54	10			1.86	2	
	48-54	10			1.86- 1.99	1	
	0-0.38	1			0.29- 0.35	1	
	0.38-	2			0.35-	2	
Clana	0.68 0.4	2			0.4	2	
Siope	0.68-0.1	3	0.79	Factor	0.4- 0.44	3	0.59
	0.1-1.39	4			0.44- 0.49	4	
	1.39-	5			0.49-	5	
	1.79	C			0.54	C	
	2.12	6			0.54- 0.59	6	
	2.12- 2.49	7			0.59- 0.63	7	
	2.49- 2.93	8			0.63- 0.68	8	
	2.93-	9			0.68-	9	
	3.37-	10			0.72	10	
	4.05				0.78		
	0.003- 0.011	1			1.94- 2.29	1	
	0.011- 0.022	2		Bifurcation	2.29- 2.73	2	
Relief	0.022-	3	0.64	ratio	2.73-	3	
Tallo	0.038-	4			3.18-	4	0.71
	0.055	5			3.54	5	
	0.073	0			3.83	U	
	0.073-	6			3.83-	6	
	0.089-	7			4.22-	7	
	0.11 0.11-	8			4.69 4.69-	8	
	0.12				5.17	•	
	0.12- 0.14	9			5.17- 5.62	9	
	0.14- 0.15	10			5.62- 6.08	10	

- **Table 4**: Rank distribution, logic behind rank distribution, correlation score and weighted
- 141 score of Relief Diversity model.

Parameters	Rank	Logic	Total correlation score(Xi)	Weighted score Xi/Ximax
Dissection index	10 rank at highly dissected area	Extreme dissected area is a very good indicator of relief diversity	4.03	0.87
Ruggedness number	10 rank at maximum rugged topography	More value of Ruggedness number enhance the relief diversity	3.57	0.77
Relative relief	10 rank at highest relative relief	Highest value of relative relief increase the relief diversity	4.63	1
Slope	10 rank at steep slope	Greater value of slope promote the relief diversity	3.65	0.79
Relief Ratio	10 rank at maximum relief ratio	More value of relief ratio demarcated more relief diversity	2.95	0.64
HG	10 rank at highest value of Hydraulic gradient	Higher value of hydraulic gradient shows greater relief diversity	2.38	0.52
Elevation	10 rank at maximum relief	Greater value of relief indicate the more relief diversity	2.66	0.58
Sinuosity	10 rank at less sinuous region	Where sinuosity is low, relief diversity is high	1.49	0.32
Asymmetry Factor	10 rank at more value of asymmetry	More asymmetry shows more relief diversity	2.71	0.59
Bifurcation ratio	10 rank at highest value of Bifurcation ratio	Higher value of Bifurcation ratio denotes the greater relief diversity	3.30	0.71

144Table 5 : Rank distribution, logic behind rank distribution, correlation score and weighted145score of Drainage Diversity.

Parameters	Rank	Logic	Total correlation score(Xi)	Weighted score Xi/Ximax
Drainage frequency	10 rank at maximum drainage frequency	Where drainage frequency is maximum Drainage diversity also maximum	5.40	0.914
Drainage density	10 rank at highest drainage density	High drainage density indicates higher drainage diversity.	5.91	1
constant channel	10 rank at more CCM value	High CCM value favorable for more	4.27	0.72

maintenance(CCM)		Drainage diversity		
Drainge intensity	10 rank at greater intensity	Maximum Drainage Intensity lead to drainage diversity	3.04	0.52
Length of over land flow(Lof)	10 rank at maximum value of Lof	High Lof indicates higher drainage diversity	5.68	0.96
6.Drainagre texture(Dt)	10 rank at maximum drainage texture	high value of DT maximized drainage diversity	5.57	0.94
7.Junction frequency	10 rank at highest junction frequency	More junction frequency positively effect on drainage diversity	3.18	0.54
Regional sinuosity	10 rank at lowest sinuosity	Drainage diversity is greater at less sinuous river zone	1.60	0.27
Infiltration number(In)	10 rank at more IN	More In value indicates more drainage diversity	4.54	0.77
Bifurcation ratio(Rb)	10 rank at highest Bifurcation ratio	Frequently flood occurrence observed in high drainage diversity area.	3.40	0.58
Form factor(Ff)	10 rank at highest form factor	Higher value of form factor highlights the absolute drainage diversity	2.82	0.48
Relative relief	10 rank at highest relative relief	High elevation represent greater drainage diversity	2.97	0.50
Ruggedness number	10 rank at greater ruggedness number	Drainage diversity is high in rough topography	5.76	0.97

148Table 6: Rank distribution, logic behind rank distribution, correlation score and weighted149score of Morphometric Diversity.

Parameters	Rank	Logic	Total correlation score (Xi)	Weighted score (Xi/Xi max)
Asymmetry	10 rank at high asymmetric value	High asymmetric value increase morphometric diversity	3.89	0.62
Drainage Density	10 rank at highest stream density	High stream density indicates higher morphometric diversity.	5.28	0.84
Stream Frequency	10 rank at maximum stream frequency	Where stream frequency is maximum morphometric diversity also maximum	5.17	0.83
Drainage intensity	10 rank at greater intensity	greater Drainage Intensity lead to more diversity	2.80	0.45
Dissection index	10 rank at higher dissected value	More dissected zone indicates more morphometric diversity	5.06	0.81
Drainage Texture	10 rank at maximum drainage texture	high value of drainage texture maximized morphometric diversity	5.05	0.81

Form Factor	10 rank at highest form factor	Higher value of form factor highlights morphometric	3.84	0.61
		variety		
Hydraulic gradient	10 rank at greater HG value	More HG value enhance the differentiation of morphometric diversity	3.07	0.49
Infiltration Number	10 rank at more IN value	More In value indicates more morphometric diversity	4.11	0.65
Junction Frequency	10 rank at highest junction frequency	More junction frequency positively effect on morphometric diversity	4.27	0.68
Bifurcation Ratio	10 rank at highest Bifurcation ratio	Frequently flood occurrence observed in high morphometric diversity area.	5.38	0.86
Relief Ratio	10 rank at highest relief ratio	High relief ratio represent greater morphometric diversity	4.05	0.65
Relative Relief	10 rank at highest relative relief	High relative relief indicates higher morphometric diversity	5.74	0.91
Ruggedness Number	10 rank at greater ruggedness number	morphometric diversity is high in rough topography	6.27	1
Slope	10 rank at steep slope	Steep slope positively control morphometric diversity	4.22	0.67
Elevation	10 rank at high elevation	Morphometric diversity influenced by high elevation	4.41	0.70



152

Fig: 2 Work Flow Diagram

153 In figure 2 detail methodological flow for constructing relief diversity model has been made. For other

154 cases same approaches have been adopted as shown here but most of the parameters are different.

155 4. RESULT AND DISCUSSION

156 4.1 Spatial data layers used for the models

A brief account of some morphometric characteristics of Kushkarani river basin are as follows: total length of the stream is 33.83 km., bifurcation ration is 4.027, length of over land low is 0.88, constant of channel maintenance is 0.568, sinuosity is 1.33, form factor of the basin is 0.268, drainage density is 1.76 km/sq km. drainage texture is 4.36, drainage intensity is 3.75, hypsometric integral (HI) is 0.48, relative relief is 93m., dissection index (DI) is 0.808, and ruggedness number is 0.062. All these values are representative to the entire basin but spatial variation is very strong.

163 Figure 3 (A-S) represent different spatial data layers used for different diversity model. Different data 164 layers behave differently towards the objective. For example in case of ruggedness number, relative 165 relief, drainage density, intensity, frequency these are high in the upper catchment of the basin 166 whereas it low in the same area i.e. form factor, bifurcation ratio and relief ratio. In the same way 167 differential behavior is also noticeable at the middle and lower catchment of the basin. For example 168 ruggedness number for this basin ranges from 0 to 0.062 (Fig. 3A). Few patches of upper and middle 169 catchments represent high range of value (0.05-0.06) covering an area of 11.6% to total area. 170 Dissection index(DI) ranges from 0.09 to 0.576 and high value is recognized at the upper catchment 171 and left portion of the middle catchment (Fig. 3B). Position of scattered hillock like landform has 172 generated this situation. Relative relief value ranges from 10-54m. (Fig. 3D) and it is high in those part 173 where dissection is high as mentioned earlier. Slope of this basin ranges from 0-4.05° and high 174 degree of slope is noticed at the upper fringe of the upper catchment (Fig. 3F). Figure 3H indicates 175 that sinuosity value ranges from 0-1.98 and no definite trend is identified over the basin as the entire 176 basin lies over the plateau fringe area. Bifurcation ratio (range: 1.93-6.07) is relatively higher in the 177 lower middle and upper part of the lower catchment (Fig. 3J). Since these areas, volume of water and 178 related functions may be energized. Drainage density of this basin ranges from 0-1.4 and it is found in 179 scattered patches over the basin (Fig. 3K).

180



181 182

(A) Ruggedness Number









(E) Relief Ratio

(F) Slope



187 188

(G) Hydraulic Gradient

(H) Sinuosity Index



189 190

(I) Asymmetry Factor

(J) Bifurcation Ratio





Fig 3 (A –S) shows the spatial data layers for constructing the models. The parameters which are used for constructing the three models are described separately in table 2 -4.

204

205 **4.2 Relief diversity model (RDM)**





207 Fig 4 (a) Weighted Relief Diversity (b) Weighted Classified Relief Diversity

208 PCA based weights composite equation for Relief Diversity

209	RDM= (Rn*0.87)+(DI*0.77)+(Rhp*1)+(S*0.79)+(Rh*0.64)+(HG*0.52)+(H*0.58)+(Si*0.32)+(Aa*0.59)+(Rb*0.71)
210	(Eq.3)

211

212 **Table 7**: Area and percentage of area under different weighted Relief Diversity classes.

Relief diversity status	WLC score	Area extent (sq.km)	% of total area
Very low	50.81 - 67.97	24.88	14.47
Low	67.97 - 79.12	38.37	22.30
Moderate	79.13 - 90.34	46.65	27.12
High	90.34- 102.40	41.80	24.30
Very High	102.40- 125.68	20.30	11.81

213

214 Relief diversity model mainly is constructed for identifying the unique geomorphological surface based 215 on relief characteristics in the river basin. Equation 3 represents the concise weighted linear 216 combination model of relief diversity. Figure 4(a) shows that very high relief diversity is found at the 217 upper catchment of the basin where ruggedness number, dissection index, elevation, relative relief 218 and slope are high. In this area, average slope is 2.02%, DI is 3.625, ruggedness number is 0.031. 219 Such conditions indicate that the landscape unit is characterized by greater soil erosion potentiality, 220 less chance of ground water recharge etc. Due to having greater potentiality of top soil loss, there will 221 be paucity of soil nutrients and therefore, agriculturally this region will not be potential enough. Tree 222 feeling incidents will be more frequent in this area. Out of the total area 11.81% area is found under 223 very high relief diversity zone (table 7). In case of very low relief diversity, it is found at the lower 224 catchment of the basin where ruggedness number, dissection index, elevation, relative relief, and 225 slope are low. Out of the total area 14.47% area are come under very low relief diversity zone. This 226 region usually will be highly potential for agricultural activities due to greater potentiality of soil

- 227 moisture and less susceptibility to soil erosion. Surface and sub surface water potentiality will also
- 228 support such economic activities.

229 **4.3 Drainage diversity model (DDM)**



230

Fig 5 (a) Weighted Composite Drainage Diversity (b) Weighted Composite Classified Drainage
 Diversity

233 PCA based weights composite equation for Drainage Diversity.

234	DDM=(Sf*0.91)+(Dd*1)+(CCM*0.72)+(Di*0.52)+(Lof*0.96)+(Dt*0.94)+(Jf*0.54)+(Si*0.27)+(In*0.77)+(Rb*0.58)+(
	Ff*0.48)+(Rhp*0.50)+(Rn*0.97)(Eq. 4)

235

236 **Table 8**: Area and percentage of area under different Weighted Drainage Diversity classes.

Drainage Diversity	WLC score	Area extent (sq.km)	% of total area
Very low	13.02 - 26.40	14.85	8.63
Low	26.40 - 35.23	45.72	26.58
Moderate	35.23 - 42.50	50.18	29.18
High	42.50- 51.09	40.14	23.34
Very High	51.09 - 68.49	21.11	12.27

237

238 Drainage diversity model has been prepared for identifying unique drainage landscape unit. Equation 239 4 depicts weighted linear composite model for drainage diversity model. Figure 5(a) and 5(b) 240 respectively illustrated the drainage diversity model in continuous and classified form. Very high 241 drainage diversity is observed at middle and upper catchment of the basin where drainage density (1-242 1.49 km/sq.km), drainage frequency (3.20-4.625 stream/sq.km), drainage texture (0.732-243 1.860/sq.km.), length of over land flow (0.340-0.713) and ruggedness number (0.048-0.0627) is high. 244 Table 8 shows that out of the total area of the basin 12.27% area falls under very high drainage diversity zone. This zone is characterized by large number of 1st and 2nd order streams and therefore 245 246 highly prone to soil loss. Constant erosion activities in this region are not also suitable for agriculture 247 activities and construction of settlement. It causes frequent felling of trees. Very low drainage diversity 248 is observed at the lower catchment of the basin where drainage diversity, drainage texture, drainage 249 frequency and ruggedness number are very low. Out of the total area 8.63% area (Table 8) comes 250 under very low drainage diversity zone. Dominance of rills and gullies accelerate the erosion process 251 and natural deforestation.

252 **4.4** Morphometric Diversity models (MDM)



254 255

253

Fig 6 (a) Weighted Morphometric Diversity

(b) Weighted classified Morphometric Diversity

256 PCA based weights composite equation for Morphometric Diversity.

MDM=(Aa*0.62)+(Dd*0.84)+(Sf*0.83)+(Di*0.45)+(DI*0.81)+(Dt*0.81)+(Ff*0.61)+(HG*0.49)+(In*0.68)+(Rb*0 .86)+(Rh*0.65)+(Rhp*0.91)+(Rn*1)+(S*0.67)+(H*0.70)....(Eq. 5)

259

257

258

260 **Table 9**: Area and percentage of area under different Morphometric diversity classes.

Morphometric Diversity	WLC score	Pixel count	Area extent (sq.km)	% of total area
1 (Lowest)	65.59-75.11	225	1.72222	1.00
2	75.11-84.64	724	5.54172	3.22
3	84.64-94.16	2593	19.84763	11.54
4	94.16-103.69	4206	32.19403	18.72
5	103.69-113.21	4983	38.14143	22.17
6	113.21-122.74	3880	29.69872	17.26
7	122.74-132.26	3220	24.64688	14.33
8	132.26-141.78	1509	11.55035	6.71
9	141.78-151.31	858	6.567398	3.82
10 (Highest)	151.31-160.84	273	2.089627	1.21
Total		22471	172	100

261

262 Incorporating all the Morphometric parameters finally diversity in Morphometric surface has been 263 constructed (Eq. 5). Figure 6 (a and b) show continuous WLC and classified WLC model for 264 morphometric diversity model respectively. Classified model highlights ten distinct morphometric 265 surfaces with distinct morphometric characters (Fig. 6b). Table 9 represents the area under different 266 morphometric diversity classes, where 1.21% area comes under very high and 1.00% area comes 267 under very low diversity surface. Highly diversified morphometric surface is geomorphological 268 picturesque but accelerated rill and gully action is majorly responsible for high rate soil erosion, fast rate tree felling etc [36]. Soil erosion rate of this area is 6-18 tons/ha/year as per the estimation of Jha 269 270 and Kapat ^[50] .Least diversified surface mainly comes under relatively depressed surface at the 271 confluence region where deposition action is more active triggered by back thrust of flow due to rise of water level in Tilpara barrage ^[51]. This surface is submerged seasonally during monsoon period. Table 272

273 10 briefly describes the statistics of different morphometric parameters in different morphometric 274 diversity zones. From this table one can assume the differences of different morphometric surface 275 units. For example, at diversity zone 10 where diversity level is high and it is explained by a number of 276 parameters like high drainage density (0.9-1.40km./sq.km.), high dissection index (0.29-0.50/sq.km.), 277 high stream frequency (2.1-4.5/ sq.km.) etc. This part experiences late youth to mature stage of the 278 cycle of erosion as indicated by hypsometric integral (0.42-0.55). On the other hand lowly diversified 279 area is characterized by low drainage density (0-0.32km./sq.km.), low dissection index (0.09-280 0.1/sq.km.), low stream frequency (0-0.5/sq.km.) etc. This zone is located at the late mature to old 281 stage of the cycle of erosion (HI=0.35-0.45).

282 Table 10: Parameter characters in different diversity clusters

Mor pho metr ic Dive rsity	Dd(<mark>km/</mark> sq.k m)	Rh p (in m)	Rn	DI	C C M	Si	Di	In	D t	L of	S(<mark>i</mark> de gr ee)	S f(st re a m n o. /s q. k m	H G	Hi	Ff	R b	Rh	H (i m)
1	0- 0.32	10- 13	0- 00 6	0.0 9- 0.1	0- 6	1- 1. 98	0. 0 5	0- 0. 5	0- 0. 5	0- 0. 2	0	0- 0. 5	1. 5- 2	0. 35 - 0. 45	0. 1- 0. 52	2- 3. 5	0- 0.0 01	62 - 65
2	0.05 - 0.46	12- 16	0.0 02- 0.0 03	0.0 9- 0.3	0- 5. 5	1. 4- 1. 9	0. 3- 1	0- 2. 5	0. 5- 0. 9	0. 1- 0. 2	0- 0. 1	0. 5- 1. 2	1. 6- 2. 9	0. 39 - 0. 42	0. 75 - 0. 89	2- 4. 2	0.0 01- 0.0 5	63 - 69
3	0.2- 1	13- 20	0.0 02- 0.0 04	0.1 - 0.4 5	3- 1 3	1- 1. 9	0. 8- 3. 5	1- 5. 5	0. 5- 1. 1	0. 1- 0. 3	0. 2- 1	0. 2- 3. 5	1. 2- 2. 1	0. 36 - 0. 46	0. 12 - 0. 56	1. 9- 5	0.0 2- 0.0 5	63 - 75
4	0.1- 1.2	12- 22	0.0 03- 0.0 06	0.0 2- 0.4 2	0- 7	0. 4- 1. 5	0. 1- 4. 5	1. 5- 5	0. 1- 1. 5	0- 0. 3	0. 5- 1. 1	0. 3- 4	0. 26 - 1. 56	0. 36 - 0. 45	0. 12 - 0. 61	2. 3- 5. 2	0.0 02- 0.1 5	75 - 83
5	0.6- 1.3	15- 25	0.0 05- 0.0 06	0.0 9- 0.3	4- 1 3. 5	0. 8- 1. 6	0. 5- 6. 5	2. 5- 6. 5	0. 4- 1. 6	0. 1- 0. 6	0. 2- 1. 3	1- 3	0. 28 - 3. 1	0. 38 - 0. 46	0. 15 - 0. 54	2. 5- 5. 5	0.0 2- 0.0 6	80 - 10 0
6	0.7- 1.35	20- 45	0.0 07- 0.0 6	0.2 - 0.4	2- 8	0- 0. 8	2- 6	1. 5- 4. 5	0. 6- 1. 4	0. 2- 0. 6	0. 5- 2. 5	0. 1- 4. 5	2- 3	0. 4- 0. 5	0. 45 - 0. 9	3- 5. 5	0.0 6- 0.0 7	78 - 11 5
7	0.8- 1.39	35- 54	0.0 1-	0.3 -	3- 8.	0. 1-	2. 5-	3. 5-	0. 7-	0. 5-	2. 5-	0. 2-	1. 5-	0. 3-	0. 2-	2. 5-	0.0 5-	10 0-

			0.0 6	0.5 6	5	1. 6	7. 1	7. 4	1. 7	0. 7 1	3. 5	3. 5	3	0. 51	0. 9	6	0.6	12 1
8	0.65 - 1.36	40- 50	0.0 2- 0.0 4	0.3 5- 0.4 6	3- 1 0	0. 2- 1. 1	4- 6	4. 5- 7. 2	0. 7- 1. 6	0. 4- 0. 6 8	3- 4	2- 4. 2	1. 0- 3. 3	0. 35 - 0. 49	0. 45 -1	3. 5- 5. 1	0.0 4- 0.0 7	11 5- 13 2
9	0.7- 1.39	30- 53	0.0 3- 0.5 2	0.3 6- 0.5 1	2- 7	0. 1- 0. 9	5- 7. 1	3- 7. 1	1. 2- 1. 8	0. 5- 0. 7	3. 5- 4. 1	2. 1- 4. 5	2. 9- 3. 33	0. 45 - 0. 51	0. 65 - 1. 0	5. 1- 5 6	0.0 6- 0.0 8	12 0- 14 0
10	0.9- 1.40	25- 54	0.0 2- 0.0 56	0.2 9- 0.5 0	2. 5- 8	0. 15 - 0. 5	4. 5- 7	3. 5- 7. 4	1. 3- 1. 7	0. 5- 0. 7 1	3. 6- 4. 1	2. 6- 4. 6	3. 0- 3. 3	0. 42 - 0. 55	0. 63 - 1. 1	5- 5. 8	0.0 6- 0.0 8	13 5- 15 5

284 **4.5** Spatial Association among the RDM, DDM and MDM

285 Correlation matrix has been carried out among three above mentioned diversity model(s) to know the 286 role of drainage and relief parameters for making the surface diversified. Table 11 shows relief 287 diversity and morphometric diversity are strongly correlated as the correlation value is 0.94 and 288 drainage diversity is positively correlated with RDM and MDM as the correlation values are 0.445 and 289 0.207 respectively. Former two cases represent significant correlation at 0.01 level of significance. In 290 an ideal river basin, relief and drainage is highly associated, elevation, slope control flow form, flow 291 density etc. but at this present case association between relief and drainage is not so strong. Relation 292 is positive (r=o.207) and also significant at 0.1 level of significance between two but not statistically 293 acceptable. Low relief diversity over this plateau fringe river is the major reason behind such low 294 spatial variability. From this correlation structure, it is clear that for bringing morphometric diversity, 295 relief parameters play stronger role than drainage diversity and drainage diversity of the landscape is 296 also controlled by relief diversity.

297

298 **Table 11**: Correlation among three models

	Relief diversity	Drainage diversity	Morphometric diversity
Relief diversity	1	0.207	0.940
Drainage diversity		1	0.445
Morphometric diversity			1

299

300 <mark>5. CONCLUSION</mark>

From the analysis of three models different unique drainage, relief and morphometric surface(s) have been recognized. Composite weighted score of the relevant parameters have been classified into equal 10 classes representing 10 morphometric surfaces. Range of value with greater magnitude of values is considered as most diversified surface with relatively dynamic processes and forms and vice versa. Each unit can be treated as unique in terms of terrain processes and forms. These are also

- 306 distinct in connection with potentiality for developing its own characteristics assemblages cultural
- 307 landscapes. Distinguished terrain processes, soil catena, soil loss, water retaining capacity, recharge,
- 308 surface water availability and runoff characters characterizes the economic activities of the study area
- 309 specially agricultural activities. So, not only for categorizing morphological units but also for devising
- 310 economic landscaping such units are important.

311 **References**

327

328

337

338

339

342

343

344

345

346

347

348

- Blankson EJ, Green BH. Use of landscape classification as an essential prerequisite to
 landscape evaluation. Landscape Urban Planning.1991;21:149-162.
- Otero PI et al. Landscape evaluation: Comparison of evaluation methods in a region of Spain.
 Journal of Environmental Management. 2007.In Press.
- Staler A N. Quantitative geomorphology of basins and channel networks. In: Chow, V.T.
 (Ed.), Handbook of applied hydrology. Mcgraw Hill Book Company, New York.1964.
- Mages NS, Jitheslal KV, Chandrasekar N. Geographical information system-based morphometric analysis of Bharathapuzha river basin, Kerala, India, Appl Water Sci .
 2013;3:467–477.
- Rastogi RA, Sharma TC. Quantitative analysis of drainage basin characteristics. Jour Soil and water Conservation in India. 1976; 26:18-25.
- Baker VR. Introduction: regional landforms analysis. In: M Short, RW Blair, Jr. (eds),
 Geomorphologyfrom space, a global overview of regional landforms. National Aeronautics
 and Space Administration, SP 486. US Government Printing Office, Washington DC. 1986:1–
 26.
 - Nautiyal MD. Morphometric analysis of a drainage basin, district Dehradun, Uttar Pradesh, J Indian Soc Remote Sens.1994; 22(4) :251–261.
- 8. Nag SK, Chakraborty S. Influence of rock types and structures in the development of
 drainage network in hard rock area, J Indian Soc Remote Sens. 2003;31(1): 25–35.
- Rudraiah M, Govindaiah S, Srinivas VS. Morphometry using remote sensing and GIS techniques in the sub-basins of Kagna river basin, Gulburga district, Karnataka, India. J
 Indian Soc Remote Sens.2008; 36: 351–360.
- Magesh NS, Jitheshlal KV, Chandrasekar N, Jini KV. GIS based morphometric evaluation of
 Chimmini and Mupily watersheds, parts of Western Ghats, Thrissur District, Kerala, India.
 Earth Sci Inform. 2012; 5(2):111–121.
 - John Wilson JS, Chandrasekar N, Magesh NS. Morphometric analysis of major subwatersheds in Aiyar and Karai Pottanar Basin, Central Tamil Nadu, India using remote sensing and GIS techniques. Bonfring Int J Indus Eng Manag Sci. 2012; 2(1):8–15.
- Van RA, Verstraeten G, Van OK, Govers G, Poesen J. Modeling mean annual sediment yield
 using a distributed approach. Earth Surf Process Landforms. 2001; 26(11):1221–1236.
 - Fernandez C, Wu JQ, McCool DK, Stockle CO. Estimating water erosion and sediment yield with GIS, RUSLE, and SEDD. J Soil Water Conserv. 2003;58(3):128–136.
 - 14. Ganapuram Sreedhara et al. Mapping of groundwater potential zones in the Musi basin using remote sensing data and GIS. Advances in Engineering Software. 2009; 40: 506-518.
 - 15. Vasanthavigar M et al. Groundwater potential zoning in Thirumanimuttar sub-basin Tamilnadu, India a GIS and remote sensing approach. Geo-spatial Information Science .2011;14(1):17 26. DOI 10.1007/s11806-011-0422-2
- 349 16. Awawdeh M et al. Integrated GIS and remote sensing for mapping groundwater potentiality in
 350 the Tulul al Ashaqif, Northeast Jordan. Arabian Journal of Geoscience. 2014;7(6):2377-2392.
- 17. Jasrotia AS, Kumar A, Singh R. Integrated remote sensing and GIS approach for delineation
 of groundwater potential zones using aquifer parameters in Devak and Rui watershed of
 Jammu and Kashmir, India. Arabian Journal of Geosciences.2016; 9: 304.

354 18. Ibrahim-Bathis K, Ahmed S. Geospatial technology for delineating groundwater potential 355 zones in Doddahalla watershed of Chitradurga district, India. The Egyptian Journal of Remote 356 Sensing and Space Science.2016; 19(2): 223-234. 357 19. Gajbhiye S., Mishra S. K., Pandey A. Prioritizing erosion-prone area through morphometric 358 analysis: an RS and GIS perspective .Appl Water Sci. 2014:4:51-61.DOI 359 10.1007/s13201-013-0129-7 360 20. Gopal KG, Saha S. Identification of soil erosion susceptible areas in Hinglo river basin, 361 Eastern India based on geo-statistics. Universal Journal of Environmental Research and 362 Technology. 2015;5(3):152-164. 363 21. Sarkar BC. A GIS approach to morphometric analysis of Damodar river basin and 364 groundwater potentiality mapping in Jharia coalfield. 2011;26 (1-4):1-70. 365 22. Avinash K, Deepika B, Jayappa K S. Basin geomorphology and drainage morphometry 366 parameters used as indicators for groundwater prospect: Insight from geographical information system (GIS) technique. Journal of Earth Science. 214:25(6): 1018-1032. 367 368 doi:10.1007/s12583-0140505-8 369 23. Nidhi K , Chowdary VM, Tiwari KN, Shinde V, Dadhwal VK. Assessment of surface water potential using morphometry and curve number-based approaches. 370 Geocarto 371 International.2016. www.tandfonline.com/doi/pdf/10.1080/10106049.2016.1195889 372 24. Dahiphale P, Singh PK, Yadav KK. Morphometric analysis of sub-basins in jaisamand 373 catchment using geographical information system. International Journal of Research in 374 Engineering & Technology.2014; 2 (6) : 189-202. 375 25. Rai PK., Mohan K., Mishra S, Ahmad A, Mishra VN. A GIS-based approach in drainage 376 morphometric analysis of Kanhar River Basin, India. Appl Water Sci.2014: 1-16. DOI 377 10.1007/s13201-014-0238-y. 26. Khatun S and Pal S. Identification of prospective surface water available zones with multi 378 379 criteria decision approach in Kushkarani river basin of Eastern India, Archives of Current 380 Research International. 2016;4(4): 1-20. DOI:10.9734/ACRI/2016/27651 381 27. Sures M. et al. Prioritization of watersheds using morphometric parameters and assessment 382 of Surface Water Potential using Remote Sensing. Journal of the Indian Society of Remote 383 Sensing.2004;32(3):249-259. 384 28. Thakkar A K., Dhiman S. D. Morphometric analysis and prioritization of mini watersheds in 385 Mohr watershed, Gujarat using remote sensing and GIS techniques. Journal of the Indian Society of Remote Sensing.2007; 35(4):321-329.DOI: 10.1007/BF02990787 386 387 29. Amani M, Safaviyan A. Sub-basins prioritization using morphometric analysis-remote sensing 388 technique and GIS-Golestan-Iran. 2015;38:56-65. International Letters of Natural Sciences 389 Online: 2015:05-06. 390 30. Rahaman SA, Ajeez SA, Aruchamy S, Jegankumar R. Prioritization of sub watershed based 391 on morphometric characteristics using fuzzy analytical hierarchy process and geographical 392 information system - A study of Kallar watershed, Tamil Nadu. Aquatic Procedia. 393 2015;4:1322-1330. 394 31. Clark JJ and Wilcock PR. Effects of land use change on channel morphology in Northeastern 395 Puerto Rico. Bulletin, Geol. Society of America. 2000;112(12):1763-1777. 396 32. Godoy, Pinto MD, Lacerda de, Drude Luiz .River-island morphological response to basin 397 land-use change within the Jaguaribe river estuary, NE Brazil. Journal of Coastal 398 Research.2014;30(2):39. 399 33. Khatun S and Pal S. Analysis of regional hypsometric integral to identify landscape evolution 400 in Kushkarani river basin. Journal of Geography, Environment and Earth Science 401 International. 2016; 6(3): 1-17. 402 403 34. Bagchi K, Mukerjee KN. Diagnostic survey of West Bengal(s), dept. of geography. Calcutta 404 University, Pantg Delta & Rarh Bengal. 1983;17(19):42-58.

405	35. Chakrabarty SC. Some consideration on the evolution of physiogrphy of Bengal, i	n
406	Chattopadhyay B. (ed.), West Bengal, Geog. Inst., Presidency College, Calcutta.1970;20.	
407	36. Pal S. Identification of Soil Erosion Vulnerable Areas in Chandrabhaga River Basin: a Mul	ti
408	criteria Decision Approach, Model. Earth Syst. Environ. 2016; 2(5):1-11. DC	וו
409	10.1007/s40808015 0052-z	
410	37. Wentworth CK. A simplified method of determining the average slope of land surfaces	3.
411	American Journal of Science. 1930;21:184-194.	
412	38. Schumn SA. Evolution of drainage systems and slopes in badlands at Perth Amboy, New	N
413	Jersey. Geol. Soc. Am. Bull.1956;67:597-646.	
414	39. Horton RE. Erosional development of streams and their drainage basins, hydrologica	٦l
415	approach to quantitative morphology. Bull. Geophys. Soc. Am. 1945; 56:275-370.	
416	10. Faniran A. The index of drainage intensity - A provisional new drainage factor. Australia	n
417	Journal of Science. 1968; 31: 328-330.	
418	1. Cox RT. Analysis of drainage-basin symmetry as a rapid technique to identify areas o	٥f
419	possible quaternary tilt-block tectonics: an example from the Mississippi Embaymen	t.
420	Geological Society of America Bulletin. 1994; 106: 571- 581.	
421	12. Sarangi A, Madramootoo C, Enright MP. Development of user interface in Arc GIS for	or
422	estimation of watershed geomorphology. The Canadian society for engineering i	n
423	agricultural. Food and Biological Sytsems. 2003:03-120.	
424	13. Miller JR. Ritter DF. Kochel RC. Morphometric assessment of lithologic controls on drainag	е
425	basin evolution in the Crawford Upland. South-Central Indiana, American Journal of	of
426	Science.1990:290:569-599. http://dx.doi.org/10.2475/ais.290.5.569	
427	14. Patton PC and Baker VR. Morphometry and floods in small drainage basins 1976;12(5); 941	-
428	952	
429	15. Strahler AN, Quantitative analysis of watershed geomorphology. Transactions, America	n
430	Geophysical Union, 1957:38 (6):913-920.	
431	16. Wu et al. Gap shape classification using landscape indices and multivariate statistics	s.
432	Scientific Report 6, 2016 doi:10.1038/srep38217	
433	7 Klinger R. Schwanghart W. Schutt B. Landscape classification using principal componer	nt
434	analysis and fuzzy classification: Archaeological sites and their natural surroundings i	n
435	Central Mongolia Geoarchaeology 2011: 3:213-233	
436	18 Carver ST Integrating multi-criteria evaluation with geographical information systems	2
430	International Journal of Geographical Information Systems 1991: 5(3):321-339	<i>.</i>
437	10 Eastman IR Idrisi for windows version 2.0: Tutorial eversions araduate school o	h
130	apography_Clark University Worcester MA: 1997	,
435	50 Jba VC Kanat S Gully erosion and its implications on land use a case study. In IHA VC	•
-++0 ///1	(Ed) Land degradation and desartification Public language Now Dolbi 2002:156 179	<i>'</i> .
441	(Lu). Lanu ucyrauaiion anu ucschinicaiion. Fubi., Jaipur anu new Deini. 2003, 130-170.	
442	or. Fai S. identification of soil erosion vulnerable areas in Chandrabhaga river basin: a mult	ו- ג
443	one one a decision approach, model. Earth Syst. Environ. 2016; 2(5):1-11; DOI 10.1007/\$40808)-
444	010-0002-2	