| ORIGINAL RE | SEARCH | PAPER |
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# DIMENSIONLESS GEOMORPHOMETRY AND DISCHARGE IN THE IKPA RIVER BASIN, NIGERIA

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#### 5 ABSTRACT

This study evaluates the relationships between dimensionless basin geomorphometry and 6 discharge in the Ikpa River. The basin was stratified into seven sub-units using [1] scheme, 7 8 Geospatial tools were used in generating data for three digital elevation modeling, while dimensionless geomorphometric parameters were generated from topographic maps (sheet 9 322 NE; sheet 322 SE; sheet 323 SW; and sheet 331 NW) of the basin area. The sampled 10 sub-basins were gauged and discharge measured by a surface float. Graphical analysis of 11 discharge revealed wide variations between months and in seasons across sub-basins with the 12 rainy season attracting highest volume of discharge and the corresponding fluvial processes. 13 14 The regression analysis yields a coefficient of multiple determination (R) of 0.937, signifying a very high effect expressed by 87.8 of the proportion of variance in dimensionless 15 geomorphometric parameters on discharge in Ikpa River Basin. Also, the computed F value 16 yields 1.439, while the Table value tested at  $(0.05)_{5/2}$  confident level offers 19.30. The result 17 led to the conclusion that variations in relief ratio, average bifurcation ratio, circularity ratio, 18 19 elongation ratio and form factor have a significant effect on discharge in Ikpa River Basin. This paper recommends for prompt installations of state of the art river gauging and 20 monitoring facilities to provide the needed information to the government, researchers, and 21 individuals for the sustainable land and watershed development options (dam and irrigation) 22 23 in the coastal plain of Ikpa River basin.

24 Keywords: River basin, geomorphometry, dimensionless parameter, discharge, coastal plain.

#### 26 **1.0 INTRODUCTION**

It is usually argued among contemporary physical geographers and earth scientists 27 28 that the emergent of landforms in a given geomorphologic unit depended on the prevalent 29 processes that acted on them over a period time and in space. For [2] perspective, the present 30 landscapes have evolved through the Tertiary and Quaternary, and often retain the imprint of 31 previous sub-aerial and marine processes. Hence relicts of depositional surfaces possess the 32 potential for geochronological interpretation through the sedimentary structures. Besides, [3] 33 observes that the definition of landforms of various sorts is an essential part of the 34 geomorphological mapping.

A river basin is an integral part of landform that can be accurately delineated using emerging geospatial technologies {Geographic Information System, Remote Sensing, and Global Positioning System} as applicable in [4,5] with its specific geomorphometric (areal, linear, relief and form) parameters determine base on mathematical and allied equations. In
fluvial geomorphology, each parameter plays important role in regulating discharge processes
in a given basin [6].

Specific geomorphometry is the measurement and analysis of specific surface features 41 42 defined by one or more processes and separated from adjacent parts of the land surface according to clear criteria of delimitation [Evans 1974 cited in 7]. The delineation of specific 43 44 geomorphometry entails drawing a closed boundary at the divide that separates the basin 45 from other basins as discussed in [3.8,9,10]. In the measurements of basin shape, an attempt 46 is made to derive dimensionless indices which consist of compactness coefficient, elongation 47 ratio, wandering ratio, fitness ratio, form factor, relief ratio, circularity ratio etc, and is useful 48 in the prediction of basin form and process relationship [11].

The approaches of drainage basin morphologic division according to [8] are grouped into three categories: The first approach emanated from the geographers' interest in regional delimitation. The second approach was concerned with the identification of the physiographic atoms out of which the matter of regions is built. These atoms were topography forming the characters of the landscape. The third basis resulted from the unitary features of both geometry and process exhibited in the basin over a period of time as recognized by Playfair [8,12] in a quantitative study of specific landform.

The concept of drainage basin has remained a focal point in modern geomorphological inquiries especially with the introduction of fluvio-geomorphological and engineering approaches to quantitative morphology. It is this topographic, hydraulic and hydrologic unity of drainage basin that provided the basis for the morphometric system of [13] as transformed and elaborated by [1,14]. The systematic description of the shape and form of a drainage basin and its sub-basin networks require specific measurements of stream network characteristics and expressing as numerical ratio or linear scale measurements [14].

Geospatial technologies are a convenient method to study the morphometric 63 64 characteristics as the satellite images provide detailed information of earth surface features 65 with its synoptic coverage, high receptivity, cost-effectiveness [4] and have been applied in 66 specific geomorphometric research [11]. But mathematical and statistical models are usually 67 more useful for the analysis and prediction of relationships between dimensionless geomorphometry (landform) and discharge (surface processes) at a basin scale. Such models 68 play a crucial role in understanding and predicting discharge and allied fluvio-69 geomorphologic hazards (erosion, flooding, and others) that have caused considerable 70

71 damages to ecosystem, environment, and man in areas within the Humid Tropics72 [15,16,17,18].

73 However, attempts have been made to established the relationship between river basin 74 discharge and dimensional geomorphometric parameters [16,19] but the influence of dimensionless variables especially relief ratio, circularity ratio, form factor, average 75 bifurcation ratio, on discharge of ungauged basins still require more attention [10,19] in 76 77 geomorphologic, geographic, and allied science literature because of their effects in 78 regulating the duration and volume flow of surficial processes within a specific landform. 79 Similarly, [7] observed that computer analysis makes the handling of multivariate indices 80 comparatively simple, and new sources of high-quality data are emerging to complement this 81 increased analytical capacity, but progress in research on specific geomorphometry may be 82 dependent on the scholars' willingness to exploit new developments, and gain appropriate 83 insights from it.

Many studies have pointed to the significant influence of basin elevation and morphology on River discharge and sediment fluxes but only a few mathematical relationships are available [20]. For process estimation, shape measurement based upon the distribution of area within the basin is probably more meaningful, and the method of [21] offers considerable potential. Besides, a greater potential may be realized if it becomes possible to recognize fluvio-geomorphologically active contributing area for a particular magnitude of discharge event, rather than employing the basin as defined by its perimeter.

91 Within this decade, several studies on the relationships of river basin morphometry 92 and discharge have been conducted using state of the art technologies such as Geographic 93 Information System (GIS), Remote Sensing and other geospatial software packages outside 94 the basin area. [22] opines trhat the estimation of various morphometric parameters can be 95 handled easily and more accurately by using GIS. However, the contributing sub-basin areas 96 often vary based on the local factors such as basin size, surface configuration, land use, 97 lithology, surface runoff, discharge volume, and others; hence infinite form and variety of 98 river basins usually respond to basic geomorphologic laws exists in nature [22].

99

#### 100 1.1 Objectives of the Study

101 The aim of this study is to evaluate the relationship between dimensionless basin 102 geomorphometric parameters and discharge in Ikpa River, Northeast of Akwa Ibom State, 103 Nigeria. To achieve aim, the following are the specific objectives: (1) to compare the

discharge characteristic of Ikpa River Basin in Akwa Ibom State; (2) to analyze the association and correlations between dimensionless geomorphometric parameters (relief ratio, average bifurcation ratio, circularity ratio, elongation ratio and form factor) and discharge in Ikpa River Basin of Akwa Ibom State.

#### 108 **1.2 Hypothesis of the Study**

109 This study is built on a null hypothesis that "the dimensionless geomorphometric 110 parameters of relief ratio, average bifurcation ratio, circularity ratio, elongation ratio, and 111 form factor have no significant effect on discharge in Ikpa River Basin, Nigeria".

#### 112 2.0 DESCRIPTION OF THE STUDY AREA

#### 113 **2.1 Location and Relief**

114 The Ikpa River is one of the  $5^{th}$  order tributaries of the Cross River. The basin is 115 located in the Northeast of Akwa Ibom State, Eastern Niger Delta of Nigeria (Figure 1). 116 Absolutely, the Ikpa River Basin is located between longitude  $7^046^{1}34.9^{11}$  and  $8^03^{1}11.9^{11}$ , East 117 of Greenwich Meridian and latitudes  $5^00^{1}3.801^{11}$  and  $5^016^{1}49.129^{11}$ , North of the Equator 118 [10,15,16,19]. The basin area covers parts of Ini, Ikono, Ibiono Ibom, Itu, Uruan and Uyo 119 Local Government Areas of the Akwa Ibom State, Nigeria.

120 Relief of the basin comprises of undulating lowland of the coastal plains which form 121 one of the eco-geomorphologic areas in the State. To [23], the terrain consists of the dissected coastal plains in the middle and Northern sections with an elevation of 100 - 150 meters a.s.l. 122 123 especially in Duem, the steep slopes of the river valley with the height of 50 - 80 meters a. s. 124 1), and a broad plain sloping gently toward the Cross River channel with elevation less than 125 50 meters a.s.l. The northern parts are traverse by undulating hills and other landforms of 126 river erosion especially valleys, gorges, ravines, meanders etc; while the downstream areas 127 are characterized by depositional landforms like terraces, floodplains, and ox-bow lakes. The 128 details of Ikpa River Basin location and elevation attributes are summarized in Figure 1 and 129 2).



130

131 Figure 1: Relief of Ikpa River Drainage Basin (Extracted from USGS DEM, 2016).

#### 132 2.2 Pedo-geomorphology and Climate

The pedo-geomorphology soils in the basin area are loose, friable, unconsolidated 133 134 ferrallitic in nature and are deficient in weatherable mineral reserves. The soils are mostly 135 deep and possess loamy sand to sandy surface especially at the up and mid-stream areas due 136 to the influence of the basin geologic formation. The basin geology is underlain by coastal 137 sands plains of tertiary and quaternary rocks of sedimentary formations. The tertiary 138 sedimentary rocks comprise mainly of the coastal plains sands which are the older tertiary 139 rocks (Benin formation) and are more prevalent at the upper and middle parts of the River 140 Basin. The quaternary rocks formed the River beds and are made of recent deposits of fluvial 141 sediment/alluvium mostly in the downstream area of Odiok Itam, Ide Uruan, Mbiakpan, 142 Afaha Nsai, and Eman Uruan Communities of Akwa Ibom State.

143 The climate the Ikpa River Basin is a tropical humid climate (Af) based on the 144 Koppen's classification system. The basin area has a mean annual rainfall of 2443.3mm with 145 double maxima [16,19,23]. The rainy season lasts between the months of April to October 146 while the dry season usually falls between the months of November to March annually. The 147 peaks of rain always occur during the months of July and September every year [23]. The 148 mean monthly temperature of the area is around 27°C with a range of about 5°C, but changes 149 do occur based on seasons. The average maximum temperature is 31°C (February) and the coldest month (July) temperature falls below  $24^{\circ}C$  [15,16]. Evaporation in the area is equally 150

- 151 high depending on the temperature. The relative humidity within the basin area is often high,
- ranging from 80 to 100 percent but basically, decrease with the increase in temperature.
- 153



- 155 Figure 2: the three Digital Elevation Model of Ikpa River Basin
- \*Projection perspective; field of view  $30^{\circ}$ ; rotation  $45^{\circ}$ ; and tilt at angle  $35^{\circ}$ .

#### 157 **3.0 METHODOLOGY AND DATA**

#### 158 **3.1 Determination of Dimensionless Geomorphometric Parameters**

159 The Ikpa River Basin was delineated using [1] ordering scheme. The basin was classed into six strata for the generation of dimensionless geomorphometric parameters and 160 161 measurement of discharge. The five dimensionless geomorphometric parameters were 162 generated using topographic maps (of Ikot Ekpene sheet 322 NE; Ikot Ekpene sheet 322 SE; 163 Uwet sheet 323 SW; and Calabar sheet 331 NW) each produced on a scale of 1:50,000 by the 164 Federal Survey Department in Nigeria. The sampled sub-basins are {Idim Duem (1), Iyere 165 Stream(2), Akpan Stream (3), Itam Stream (4), Amoor Stream (5), Iba Oku Stream (6), and 166 Ufak Efion (7) as depicted in figure 1 and 2. The mathematical equations for computing each 167 of the dimensionless geomorphometric parameter are summarized on Table 1.

168

| Dimensionless          | Mathematical Formula (Equation)                                      | Reference  |  |
|------------------------|--|------------|--|
| Parameter              |  |            |  |
| Bifurcation Ratio      | $Rb = N\mu/N_{\mu+1}$ . Where, $Rb = Bifurcation ratio; N\mu$        | [1,14]     |  |
|                        | = No. of stream segments of a given order and                        |            |  |
|                        | $N_{\mu+1}$ = No. of stream segments of next higher order.           |            |  |
| Mean Bifurcation Ratio | Rba = Average bifurcation ratios of all orders.                      | [1,26]     |  |
| Elongation Ratio       | Re= $\sqrt{A/\pi/Lb}$ = 1.128A <sup>0.5</sup> /Lb; Where, A= Area of | [10,24]    |  |
|                        | the basin; Lb = (Maximum) Basin length. $\pi = 3.14$                 |            |  |
| Circularity Ratio      | $Rc = 4\pi A/P^2$ ; = 12.57A/P <sup>2</sup> ; Where, A = Basin area  | [1,30]     |  |
|                        | $(Km^2)$ and P = Basin Perimeter.                                    |            |  |
| Form Factor            | $Rf = A/Lb^2$ ; Where A = Area of the basin and Lb =                 | [13,24,29] |  |
|                        | (Maximum) basin length.  |            |  |
| Relief Ratio (Rr)      | Rr = H / Lb Where, $H = basin relief (m) and Lb =$                   | [24]       |  |
|                        | Basin length (m).  |            |  |

#### **169** Table 1: Selected Dimensionless Specific Geomorphometric Parameters.

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#### 171 **3.2 Determination of Discharge**

In lkpa River Basin, regular observations and records of discharge were taken weekly and their mean determine for each month making a total of six months (three months each for the dry season and the rainy season) from the sampled sub-basins (see appendix 1). This was done starting from the onset, through the middle and toward the end of each season using instruments such as tape/ranging poles, graduated steel band and their velocity over time in cubic meters per second (m<sup>3</sup>sec) using stopwatch (see Agor, 2008; Umo, 2014; Umo et al, 2018).

179 The procedures adopted for gauging the each sub-basin include: - division of each stream into five segments; followed by the determination of the cross-sectional area (depth 180 multiplied by cross-channel bank) for each of the sampled streams at a distant of 12 meters 181 182 apart, where the stream channel was relatively straight and free from obstruction on meander belt and others on the bank [16,25]. The velocity was measured by means of surface float 183 184 using orange, for the three segments while the flow velocity was taken and the mean 185 multiplied by 0.85 to overcome errors emanating from the effects of wind and cross-currents as recently emphasized in [19,26]. The formulas are expressed as follows: 186

187 Discharge 
$$(Q) = AV$$
. Where  $A = Cross$ -sectional Area;  $V = Velocity$ 

188 Cross-sectional Area = Total Stream Segment (depth)  $\{a + b + c + d + e\}$ 

189

Total Stream Segment

190 Where; a, b, c, d and e = average depths of the different segments.

191 XY = Total width of the stream at the point of measurement. Velocity (V) = Flow
192 Distant/Time.

#### 193 **3.3 Data Analysis and Hypothesis Testing**

Discharge characteristic in the basin area was assessed descriptively using graphs. A multiple linear regression model was used to examine the effect of the five dimensionless geomorphometric parameters on discharge. ANOVA was used to test for significance of the combined effect of the five geomorphometric parameters on discharge variation. Correlation model was employed to assess the influence of individual geomorphometric parameter on discharge in the Ikpa River. The formula for linear regression model is expressed as follows:

200  $Y = a + b_1 X_1 + b_2 A_2 + b_3 X_3 + b_4 X_4 + b_5 X_5 + e$  Where Y = Discharge; a = Constant value; e =201 Standard Error of the estimate.  $b_1$  to  $b_5 =$  beta coefficients:  $X_1$  to  $X_5 =$  Geomorphometric 202 parameters.

The rationale for the choice of multiple regression model in analyzing the influence of the five dimensionless geomorphometric parameters on discharge is based on the fact that, it is often used to account for variations on dependent variable (discharge) on the linear combination of independent variables (morphometric parameters); estimate errors associated with a model; and to generate an equation which provides estimate of one variable on the others [see 27]

#### 209 4.0 RESULTS AND DISCUSSION

#### **4.1 Seasonal Variations of Discharge in Ikpa River**

The periodic measurements of discharge within the sampled sub-basins in the Ikpa River are shown in appendix 1. The Summaries for mean monthly and seasonal characteristics of discharge are presented as Table 2 and figure 3 in this section for comparative purposes. In figure 3, the monthly and average discharge reflect high variation in discharge amount between seasons with much proportion of flow occurring during the rainy season due to the intense rainfall events while dry season discharge volume tend to fall due to the influence of high temperature and evaporation in the basin area.

218 Figure 3: Summary of Seasonal Discharge Characteristics in Ikpa River Basin (m<sup>3</sup>/sec)



Similarly, variations in discharge characteristics are also observed across the sampled 220 221 sub-basins in each month and in season. The Duem stream contributes the highest percentage of 50.4, followed by Iyere stream contributing 19.6 percent, thus making a total of 70 percent 222 of the seasonal discharge as reflected in Table 2. The remaining sub-basins (Akpan, Itam,) 223 224 accounted for only 30 percent of seasonal discharge in the Ikpa River basin. Similar finding 225 has been reached by [15] in his analysis of rainfall pattern and runoff regime of Iba Oku 226 River. However, the differences if that his study focused on longtime trend using time series, 227 variance, and secondary data from the University of Uyo meteorological station.

| Sub –Basin<br>(stream) | Dry<br>Season | Percent | Rainy<br>Season | percent | Mean Total<br>Discharge | Average<br>Discharge | Percent (%) |
|------------------------|---------------|---------|-----------------|---------|-------------------------|----------------------|-------------|
| Idim Duem              | 39.97         | 49.68   | 89.78           | 50.80   | 129.75                  | 64.88                | 50.4        |
| Iyere                  | 14.83         | 18.43   | 35.60           | 20.14   | 50.43                   | 25.22                | 19.6        |
| Akpan                  | 9.01          | 11.20   | 15.79           | 8.93    | 24.80                   | 12.40                | 9.7         |
| Itam                   | 0.75          | 0.93    | 1.33            | 0.75    | 2.08                    | 1.04                 | 0.8         |
| Amoor                  | 10.50         | 13.05   | 24.01           | 13.58   | 34.51                   | 17.26                | 13.4        |
| Iba Oku                | 5.39          | 6.70    | 10.25           | 5.80    | 15.64                   | 7.82                 | 6.1         |
| Total                  | 80.45         | 100     | 176.76          | 100     | 257.21                  | 128.61               | 100         |

Table 2: Mean Discharge Variations in Ikpa River Basin  $(m^3/sec)$ .

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#### 230 4.2 Analysis of the Effect of Dimensionless Basin Geomorphometry and Discharge

The regression analysis of the multiple effects of relief ratio, average bifurcation ratio, circularity ratio, elongation ratio and form factor on discharge variation is presented on Table 3. The model result yields a coefficient of multiple determination (R) of 0.937 which is a very strong positive effect of the five geomorphometric variables on discharge in Ikpa River Basin. This is more evident as the R square of 0.878 indicates that 87.8 percent of the proportion of variance in discharge is attributed to the five independent variables while the remaining 11.2 percent could be accounted for by others factors like climate, vegetation, land use, geology and others that were not considered in this present study. Similarly, the adjusted R square of 0.268 suggests that 26.8 percent of the total variation in discharge is influenced by the five independent variables alone. The standard error of estimate associated with the model is 17.813.

Table 3: Summary of the Regression Model

| - |       |       |                   |                   |                        |  |  |  |
|---|-------|-------|-------------------|-------------------|------------------------|--|--|--|
|   | Model | R     | R Square          | Adjusted R Square | Std. Error of estimate |  |  |  |
|   | 1     | 0.937 | 0.878             | 0.268             | 17.813                 |  |  |  |
| 2 | D 11  |       | <b>D</b> · · · D' |                   |                        |  |  |  |

243 Predictors: (Constant), Form\_Factor, Ave\_Bifurc\_Ratio, Circularity\_Ratio, Elongation\_Ratio, Relief\_Ratio.

244 The Analysis of Variance was employed to test for a significant effect of the five 245 dimensionless geomorphometric parameters on discharge in Ikpa River Basin. The result of 246 the model presented on Table 4 reveals that the sum of squares associated with the model is 247 2283.775; the sum of squares associated with the residual yields 317.317, and the total sum of 248 squares associated with the model offers 2601.092. The computed F value yields 1.439, while 249 the Table value tested at  $(0.05)_{5/2}$  confidence level offers 19.30. Therefore, the null 250 hypothesis is accepted and inferred that variations in relief ratio, average bifurcation ratio, 251 circularity ratio, elongation ratio, and form factor have no significant effect on discharge 252 regime in Ikpa River Basin. This finding is contrary to [19] whose study of the basin 253 dimensional attributes using ANOVA model led to the conclusion that discharge is not 254 influenced by the dimensional parameters in the same.

Table 4: ANOVA Model of Basin Morphometry and Discharge

| Mo | odel       | Sum of Squares | Df | Mean Square | F     |
|----|------------|----------------|----|-------------|-------|
| 1  | Regression | 2283.775       | 5  | 456.755     | 1.439 |
|    | Residual   | 317.317        | 2  | 317.317     |       |
|    | Total      | 2601.092       | 7  |             |       |

Predictors: (Constant), Form\_Factor, Ave\_Bifurc\_Ratio, Circularity\_Ratio, Elongation\_Ratio, Relief\_Ratio.

From Table 5, a partial regression coefficient is employed to assess the effect of each geomorphometric variables on discharge in Ikpa River Basin. The unstandardized coefficients of B yields as following: constant value (-60.658), relief ratio (-81.468), average bifurcation ratio (37.565), circularity ratio (8.263), elongation ratio (0.158), and form factor (-8.648). Similarly, the result of standardized coefficients for selected variables reveals that relief ratio is -463, average bifurcation ratio is 0.857, circularity ratio is 0.055, elongation 261 ratio is 0.010 and form factor is -0.140. Based on these standardized beta coefficients, the 262 multiple linear regression is modelled in the equation as thus:

263  $Y = -60.658 - 0.453x_1 + 0.857x_2 + 0.055x_3 + 0.010x_4 - 0.140x_5 + e(17.813)$ 

264 Where  $x_1$  is relief ratio;  $x_2$  is average bifurcation ratio;  $x_3$  is circularity ratio;  $x_4$  is elongation 265 ratio and  $x_5$  is form factor while e is the standard error of the estimate.

266 Table 5 shows the measurement of the strength of relationship of each 267 geomorphometric parameter on discharge when others are held constant. First, only 268 elongation ratio is significant at 0.05 confidence level. Secondly, the partial correlation 269 coefficients for zero order reveals that average bifurcation ratio (0.850) and form factor 270 (0.056) exercise positive influence while elongation ratio (-0.487) and relief ratio (-0.332) 271 and circularity ratio (-0.117) exercised the negative influence on discharge variation. 272 Therefore, the positive correlation suggest that an increase in a given unit of any of the 273 independent variables while holding others constant will likely attracts a corresponding increase in the unit of discharge and vice versa; while the negative correlation is an indication 274 275 that an increase in a given independent variable by certain unit will likely lead to a decline in 276 the dependent variable and vice versa.

| 277 | Table 5: Regression and Correlation Coefficients |  |
|-----|--|--|
|     |  |  |

| Model               | Unstandardized<br>Coefficients |            | Standardized<br>Coeff. |        |       | Correlations |         |
|---------------------|--------------------------------|------------|------------------------|--------|-------|--------------|---------|
|                     |                                |            |                        | t      | Sign. | Zero         | Partial |
|                     | В                              | Std. Error | Beta                   |        |       | order        |         |
| 1(constant)         | -60.658                        | 105.499    |                        | -0.575 | 0.668 |              |         |
| Relief Ratio        | -81.486                        | 157.077    | -0.463                 | -0.519 | 0.695 | -0.332       | -0.460  |
| Average bifurcation | 35.565                         | 19.365     | 0.857                  | 1.837  | 0.317 | 0.850        | 0.878   |
| Circularity Ratio   | 8.371                          | 137.936    | 0.055                  | 0.061  | 0.961 | -0.117       | 0.061   |
| Elongation Ratio    | 0.158                          | 8.263      | 0.010                  | 0.019  | 0.988 | -0.487       | 0.019   |
| Form Factor         | -8.648                         | 25.639     | -0.140                 | -0.337 | 0.793 | 0.056        | -0.320  |

#### 278 **5. SUMMARY OF FINDINGS**

279 The descriptive analysis of discharge depicted in figure 3 and table 2 reveal 280 remarkable variations in discharge characteristics across the sampled sub-basins in each 281 month and in season. The Duem stream contributes the highest percentage of 50.4, followed 282 by Iyere stream contributing 19.6 percent, thus making a total of 70 percent of the seasonal 283 discharge as reflected in Table 2. The remaining sub-basin accounted for only 30 percent of 284 seasonal discharge in the Ikpa River basin. This result suggests that the flow regime and 285 fluvial processes (flood, sediment yield etc) are most prevalent in the Duem and Iyere 286 streams.

287 From hypothesis one, the effects of five dimensionless geomorphometric parameters 288 (relief ratio, average bifurcation ratio, circularity ratio, elongation ratio and form factor) on 289 discharge in Ikpa River Basin was assessed using multiple linear regression model. The 290 model summary reveals a very strong positive multiple coefficient (0.937) which represents 291 87.8% of the proportion of variance explained by the five dimensionless geomorphometric 292 variables alone. A test of significance using ANOVA gave an F-value of 1.439, which is less 293 than the Table value of 19.30. Surprisingly, the null hypothesis is accepted. It is concluded that variations in relief ratio, average bifurcation ratio, circularity ratio, elongation ratio and 294 295 form factor have no significant effect on discharge in Ikpa River Basin. Thus, attesting to 296 [19] observations that discharge in the Ikpa River is mostly control by dimensional 297 parameters.

298 The implication is that other factors such as the geologic formation, vegetation, land use, rainfall, dimensional geomorphometry. This finding is surprising considering the high 299 300 combined effect of the regression model; hence it tends to debunk [6] observation at Maros 301 River that basin shapes and relief parameters usually provide a better understanding of the 302 fluvial processes because the ways in which floods are formed depend on it. The 303 contradictory results could relate to the differences in basin size, level of local 304 geomorphology, climate, analytical tools, and ideas. The linear regression equation of the 305 effect of each dimensionless geomorphometric parameter was modelled as follows: Y = -306  $60.658 - 0.463x_1 + 0.857x_2 + 0.055x_3 + 0.010x_4 - 0.140x_5$ . From the linearized equation, Relief 307 ratio is x1, average bifurcation ratio is x2, circularity ratio is x3, elongation ratio is x4 and form 308 factor is x<sub>5</sub>.

The positive effect implies that every increase in an independent variable by a given unit while holding others constant, there will lead to a corresponding rise in the basin discharge by a given unit. Similarly, the result of the negative partial regression implies that decrease in a geomorphometric variable by a given unit while other factors remain constant will lead to an increase in discharge by a given unit and vice versa. These findings suggest that fluvial response of watersheds changes in response to spatial variations in the geomorphometric variables.

The elevation models presented as figure 2, affirmed [23] observation that the prevalence of gullies, ravines and sheet and other geomorphologic hazards on the upstream areas of Obotme, Ididep, Ikpa, Iba Oku, Nsan and others are due to the interactions between

the high relief and discharges over time-scale. Similarly, flash flood events Okpoto-Ididep, Ide-Uruan and Ufak Effion are associated with low relief. Indeed, recent studies on the Ikpa drainage basin morphology conducted by [10,15,16] attested that the Ikpa River Basin is the product of normal erosional and climatic processes.

#### 323 6. CONCLUSION

There are clear indications from the findings that a strong unity exists between dimensionless geomorphometric parameters and mean discharge in Ikpa River Basin. In the upstream and middle stream areas, elevations range from high to moderate (figure 2) and the rivers flow velocities were high and with sub-basins possessing more energy to do work especially erosion (hydraulic action, abrasion, attrition) and transportation which together trigger gully, channel migration and sliding activities on the coastal sedimentary deposits.

Within the down-stream where the discharge volume is high with low energy to transport loads, deposition is more prevalent and this invariably encourages sediment yield, silting of the channel, creation of ox-bow lakes, and flooding. It is concluded that any alteration on the pre-existing form-process relationships will likely attract a chain of reactions which may manifest in form of geomorphic and hydrologic disasters like erosion, flooding, land sliding and others because of the homogenous climatic and geologic formations in the basin area.

337 To guide against the prolong negative geomorphic and hydrologic responses as affirmed in various reports [10,16,19,23,28] regarding the Ikpa River basin, this paper 338 339 recommend for prompt installations of state-of-the-art river gauging and monitoring facilities 340 especially at Duem and Iyere streams that control a total of 70 percent of the seasonal 341 discharge and allied activities in the Ikpa River basin area (see Figure 3). Adequate data on 342 fluvial discharge will provide appropriate information that will be crucial to the government, 343 individual, geomorphologists, hydrologists, engineers, agriculturalists, and other scientists on how to enhance sustainable land and watershed development (e.g. dam/irrigation) 344 345 construction, and erosion control) in the Ikpa River basin. Such capital projects will boost 346 crop productivities of the people in the area.

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| 443 | <b>APPENDIX ONE:</b> Summary of Weekly/monthly/Seasonal Basin Discharge (m <sup>3</sup> s <sup>-1</sup> ) |          |          |       |       |       |           |  |
|-----|---|----------|----------|-------|-------|-------|-----------|--|
|     | Sub-basins  | December | February | March | June  | July  | September |  |
|     | 1 (week 1)  | 14.57    | 10.25    | 8.90  | 17.84 | 21.42 | 26.50     |  |
|     | 2   | 3.50     | 5.21     | 2.29  | 4.99  | 6.93  | 12.25     |  |
|     | 3   | 2.78     | 2.48     | 2.34  | 3.28  | 3.86  | 4.55      |  |
|     | 4   | 0.66     | 0.13     | 0.12  | 0.28  | 0.22  | 0.50      |  |
|     | 5   | 2.83     | 2.21     | 1.76  | 5.06  | 6.18  | 5.75      |  |
|     | 6   | 1.32     | 1.33     | 1.22  | 1.65  | 2.21  | 3.30      |  |
|     | 1 (Week 2)  | 15.96    | 10.29    | 10.41 | 19.61 | 18.87 | 29.27     |  |
|     | 2   | 3.62     | 4.77     | 2.75  | 5.76  | 8.13  | 12.21     |  |
|     | 3   | 2.81     | 2.15     | 2.38  | 2.17  | 4.07  | 4.59      |  |
|     | 4   | 0.15     | 0.14     | 0.12  | 0.26  | 0.34  | 0.33      |  |
|     | 5   | 3.05     | 2.58     | 2.46  | 5.37  | 5.97  | 6.71      |  |
|     | 6   | 1.41     | 1.34     | 1.25  | 1.60  | 2.26  | 3.57      |  |
|     | 1 (Week 3)  | 14.63    | 9.96     | 12.13 | 21.40 | 18.89 | 23.29     |  |
|     | 2   | 4.12     | 4.39     | 3.09  | 6.40  | 7.36  | 10.40     |  |
|     | 3   | 3.15     | 2.26     | 1.60  | 2.61  | 4.74  | 4.94      |  |
|     | 4   | 0.18     | 0.11     | 0.15  | 0.26  | 0.35  | 0.38      |  |
|     | 5   | 2.74     | 2.45     | 3.37  | 6.28  | 6.18  | 6.74      |  |
|     | 6   | 1.35     | 1.44     | 1.43  | 1.66  | 2.48  | 3.77      |  |
|     | 1 (Week 4)  | 14.49    | 8.39     | 10.69 | 27.72 | 20.21 | 24.31     |  |
|     | 2   | 4.04     | 3.33     | 3.36  | 7.16  | 9.47  | 15.75     |  |
|     | 3   | 1.12     | 2.16     | 1.79  | 3.36  | 4.22  | 4.97      |  |
|     | 4   | 0.20     | 0.12     | 0.16  | 0.27  | 0.50  | 0.31      |  |
|     | 5   | 2.84     | 2.60     | 2.61  | 5.45  | 6.03  | 6.17      |  |
|     | 6   | 1.50     | 1.30     | 1.28  | 2.07  | 2.98  | 3.21      |  |